

The planetary environment - II (Earth's ionosphere)

Lecture 6

W.K. Tobiska

Space Environment Technologies

Announcements

Contributions

daily space news and events from JPL

- <http://www2.jpl.nasa.gov/calendar/>

tutorial on solar wind and interplanetary magnetic field

- http://www-ssc.igpp.ucla.edu/ssc/tutorial/solwind_magsphere_tutorial.pdf

“Solar Wind Flows From Magnetic Funnels on the Sun”

- http://www.universetoday.com/am/publish/solar_wind_source.html?2242005

Bastille Day event

An article on the Bastille Day 2000 Event we covered in class.

http://science.nasa.gov/headlines/y2000/ast14jul_2m.htm The article brings together solar flares, sun's rotation, and auroras

Contributions

Contributions

CME and comet collision from STEREO

a "collision" between a CME and a comet

http://www.nasa.gov/mission_pages/stereo/news/encke.html

Mars Exploration Rover video clip

mission of the Mars Exploration Rover

<http://www.youtube.com/watch?v=eKGAvtRgzTw>

Aurora over Canada

aurora in British Columbia, Canada, 2006

<http://www.youtube.com/watch?v=qlXs6Sh0DKs>

Solar neutrinos and flares affect radioactive decay rates?

<http://www.symmetrismagazine.org/breaking/2010/08/23/the-strange-case-of-solar-flares-and-radioactive-elements/>

Contributions

CME can change direction

http://science.nasa.gov/science-news/science-at-nasa/2010/21sep_zigzag

Analysis of April 2002 storm event

<http://cires.colorado.edu/events/rendezvous/2007/posters/l18l.pdf2>

Collapse of thermosphere

<http://hubpages.com/hub/The-Thermosphere-Layer-of-Earths-Atmosphere-is-Collapsing>

Air Force satellite to monitor space debris:

<http://bigthink.com/ideas/24204>

Solar probe has 2018 launch

<http://www.spaceflightnow.com/news/n1009/03solarprobe/>

Solar wind interaction with airless Solar System bodies

<http://www.sciencedaily.com/releases/2010/09/100923190805.htm>

Contributions

Contributions- GCRs affect servers

- Sun Microsystems Enterprise flagship servers were experiencing mysterious crashes in 2000. Turns out their L2 caches had defective error protection and neutron strikes (due to GCRs) in these areas would cause the servers to crash. This affected Bell systems, America Online, Ebay, etc. Turns out Verisign moved to IBM Unix servers because of it. So space weather is not only an Astro issue!

Contributions

Contributions- Quantas flight SEE

- Quantas flight in 2006 had air data inertial reference unit single event effect failure traced most likely due to GCR radiation
- Sudden drop in altitude; 122 injured, 12 serious, 39 hospitalized



Contributions

Contributions

Photos from the stratosphere for less than \$2000

– Article:

<http://www.newscientist.com/blog/technology/2006/09/spaceflight-on-cheap.html>

– Photos:

<http://www.srcf.ucam.org/%7Ecuspaceflight/nova1selected/index.html>

Miscellaneous

- **Book on “Space Weather & Telecommunications” by John M. Goodman (see syllabus on DEN site)**

ISS from STS 115 and Earth



Lecture Overview

The planetary space environment - II

Earth

Ionospheric physics

Thermospheric coupling

Electron-ion production and loss processes (solar
EUV photoionization, charged particle
precipitation, Joule heating, waves, winds)

Ionospheric Structure and Dynamics

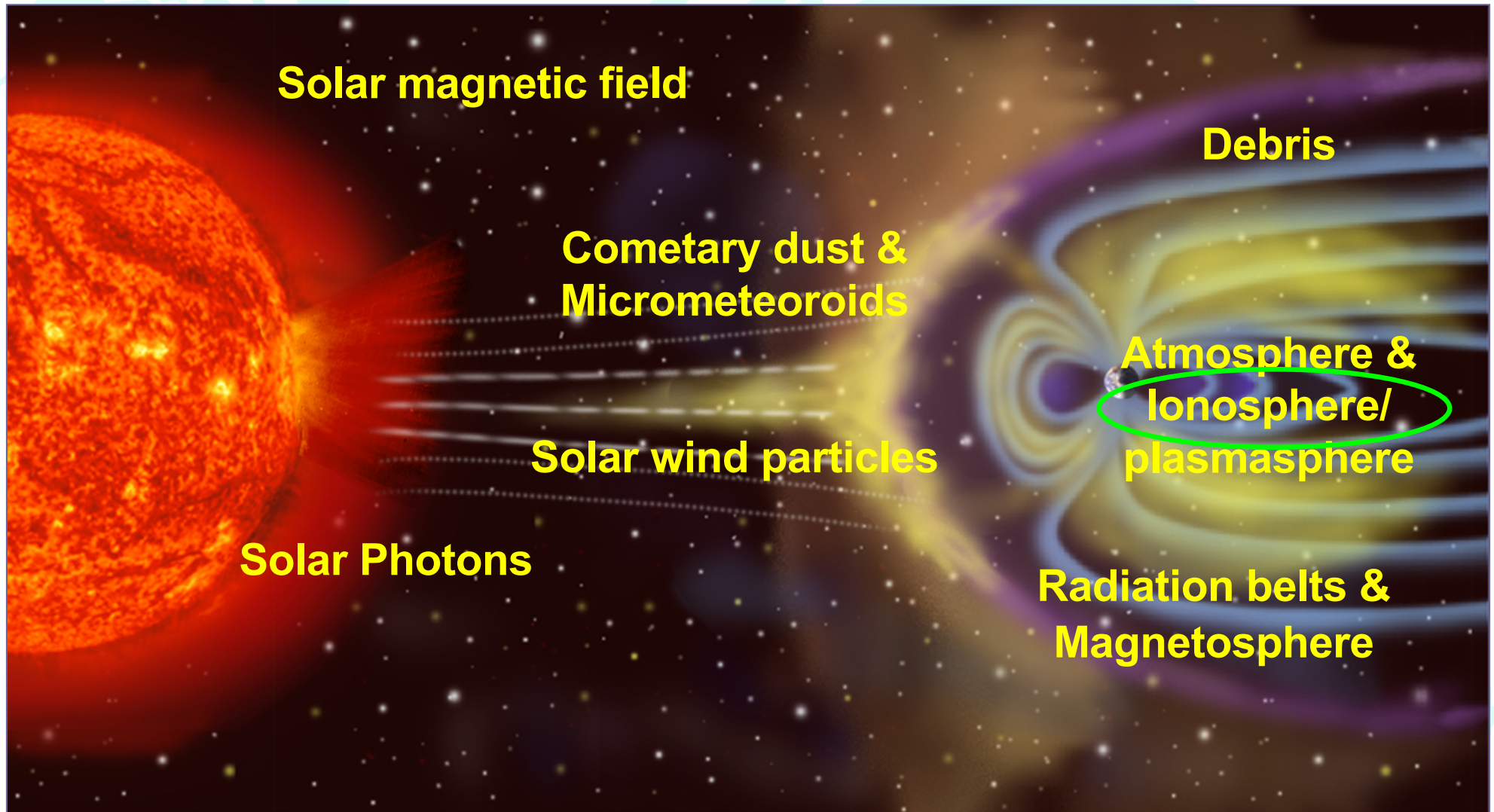
Ionospheric features

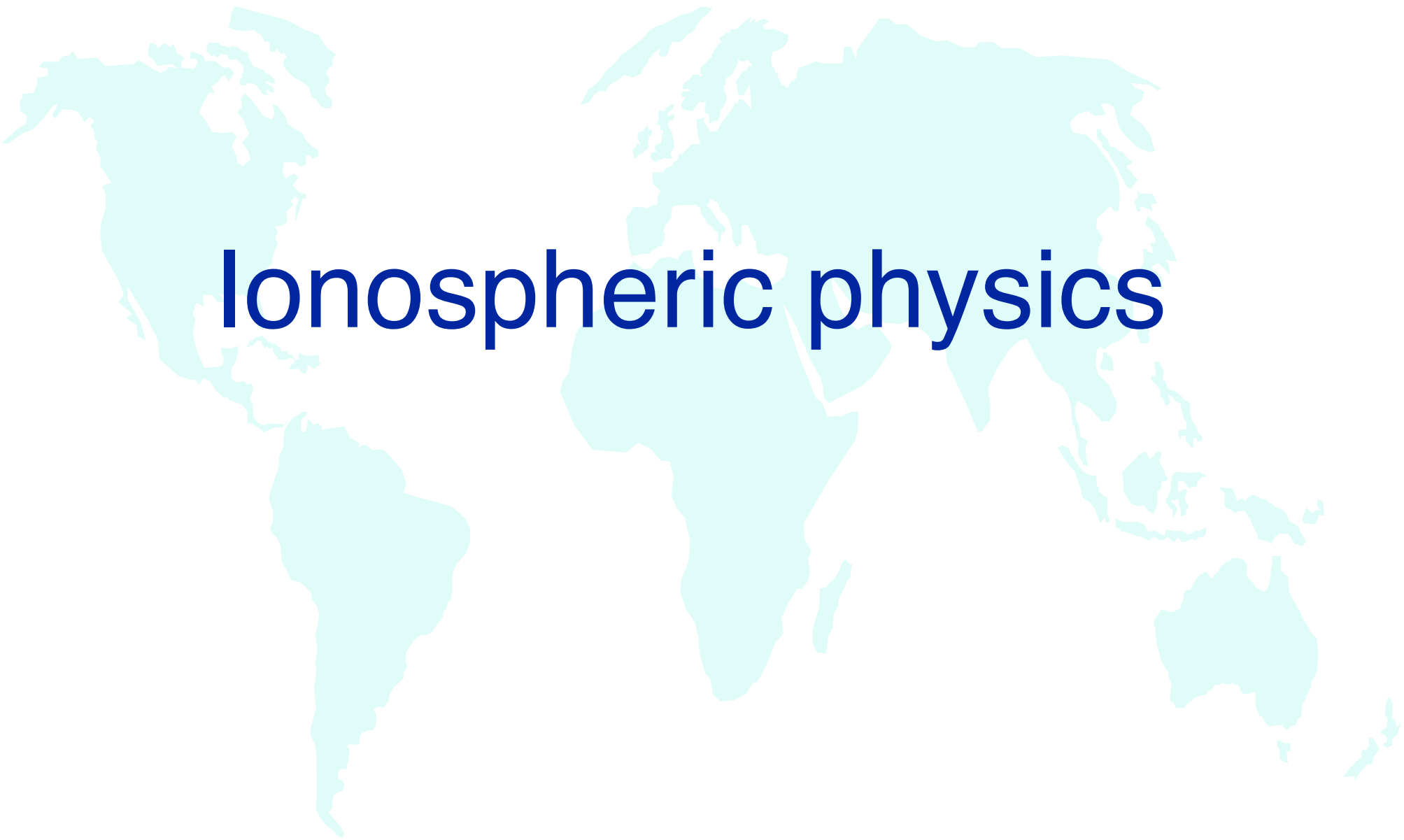
Homework

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The planetary space environment

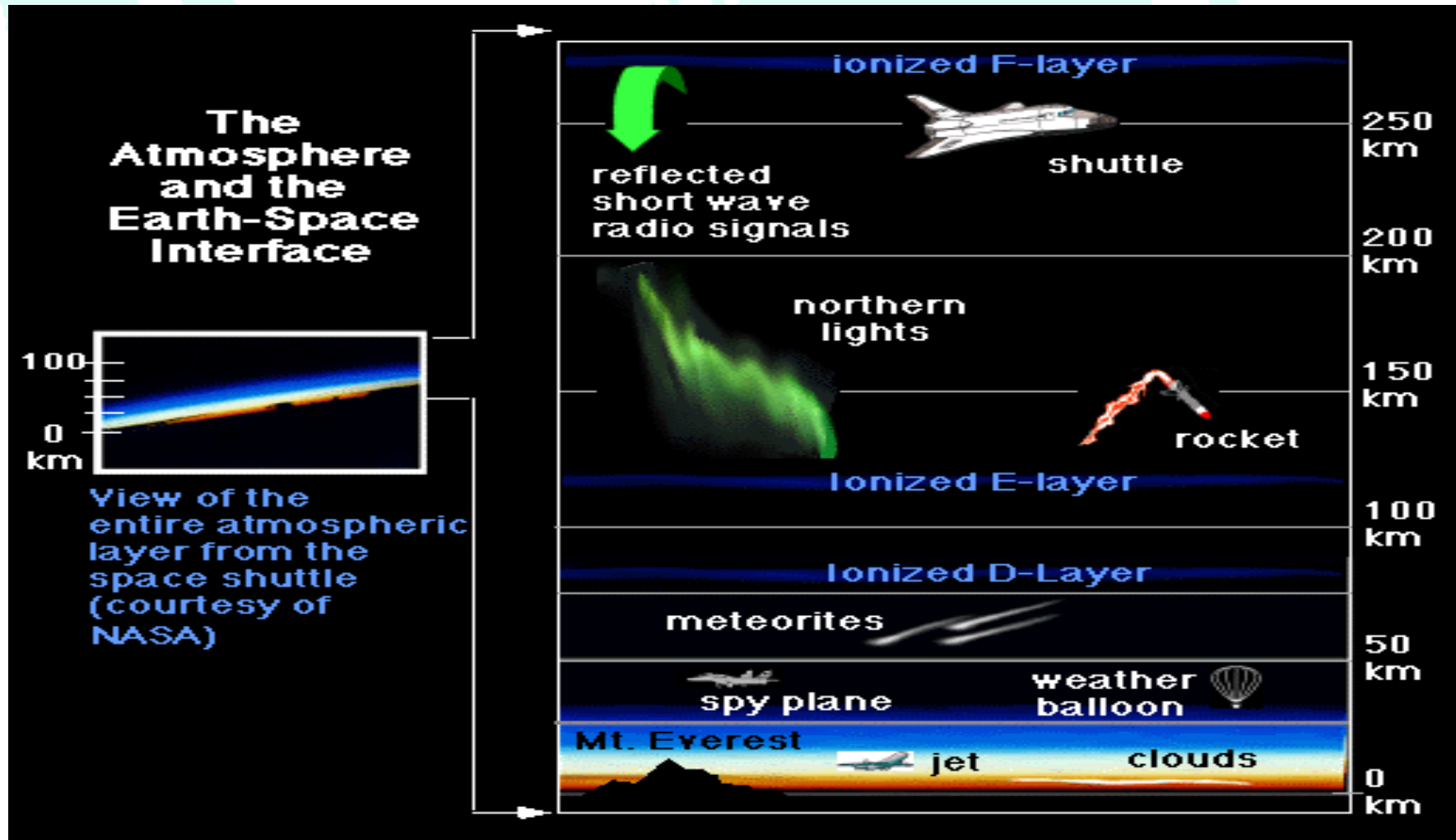
The space environment



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Ionospheric physics

The Ionosphere

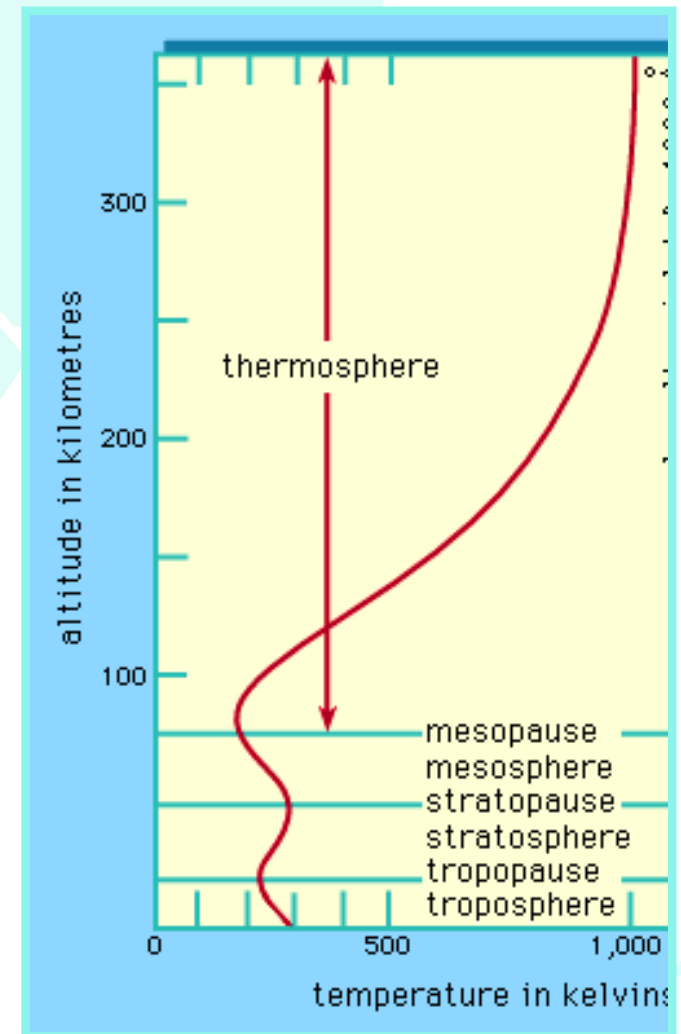


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Thermosphere coupling

The thermosphere

- ◆ Atmosphere heated by solar EUV
- ◆ Temperature increases with altitude to exobase (~ 500 km)
- ◆ Below 120 km (turbopause), atmosphere is turbulently mixed
- ◆ Above 120 km, atmosphere is stratified into layers of atomic species by molecular diffusion
- ◆ Thermosphere is electrically **NEUTRAL**



First Principles

1. Equation of Energy (Conservation of Energy):

Heat Production + Heat Loss + Heat Transport = 0

$$\frac{\partial n\varepsilon}{\partial t} + \nabla \cdot (n\varepsilon \mathbf{v}_0) + \nabla \cdot \mathbf{E} + p \nabla \cdot \mathbf{v}_0 - \sum n_i \mathbf{X}_i \mathbf{V}_i = P - L \quad (6-1a)$$

$$\mathbf{E} = -\lambda \nabla T + T \sum_i c_{pi} \rho_i \dot{\mathbf{V}}_i \quad (6-1b)$$

for n =particle concentration, $\mathbf{E}$ = heat flux, ε = total translational energy, and $\mathbf{X}_i$ is any external force acting upon the particles

2. Equation of Continuity (Conservation of Mass):

Particle Production + Particle Loss + Particle Transport = 0

$$\frac{\partial n}{\partial t} + \nabla \cdot (n \mathbf{v}_0) + \nabla \cdot \left(\sum_i n_i \mathbf{V}_i \right) = 0 \quad (6-2)$$

for $\mathbf{V}_i$ = diffusion velocity and $\mathbf{v}_0$ = mass average velocity

Credit: Tobiska

First Principles (con't.)

3. Equation of Motion (Newton's 2nd Law):

Force = mass $\times$ acceleration

$$F = ma = m \frac{dU}{dt} \quad (6-3)$$

4. Equation of State (Ideal Gas Law):

pressure $\approx$ concentration $\times$ temperature

$$p = nkT \quad (6-4)$$

Equation of Energy

Rate of change of energy in
single packet of gas is given by

$$\frac{\partial \varepsilon}{\partial t} = \text{Heat production rate} + \\ \text{Heat loss rate} + \\ \text{Heat transport rate}$$

Equation of Energy

We take a single packet of gas and examine its energy density which can be written as

$$\epsilon = \underbrace{\overbrace{c_P}^{\text{specific heat at constant pressure}} T}_{\text{internal energy}} + \underbrace{\frac{V^2}{2}}_{\text{kinetic energy}} \quad (6-5)$$

Equation of Energy

Heat production: solar irradiances

- **primary global heating** from absorption of solar UV/EUV/XUV radiation on the dayside - 2 processes

photodissociation: e.g. $\text{O}_2 + h\nu \rightarrow \text{O} + \text{O}$

ionization: e.g. $\text{NO} + h\nu \rightarrow \text{NO}^+ + e^-$

- **secondary local heating** from excess energy (kinetic energy)

Joule heating, electron precipitation, thermal conduction and vibrational, rotational, elastic collisions via eddies and turbulence

Credit: Tobiska

Ionizing radiation sources

Solar radiation - primary global ionization source

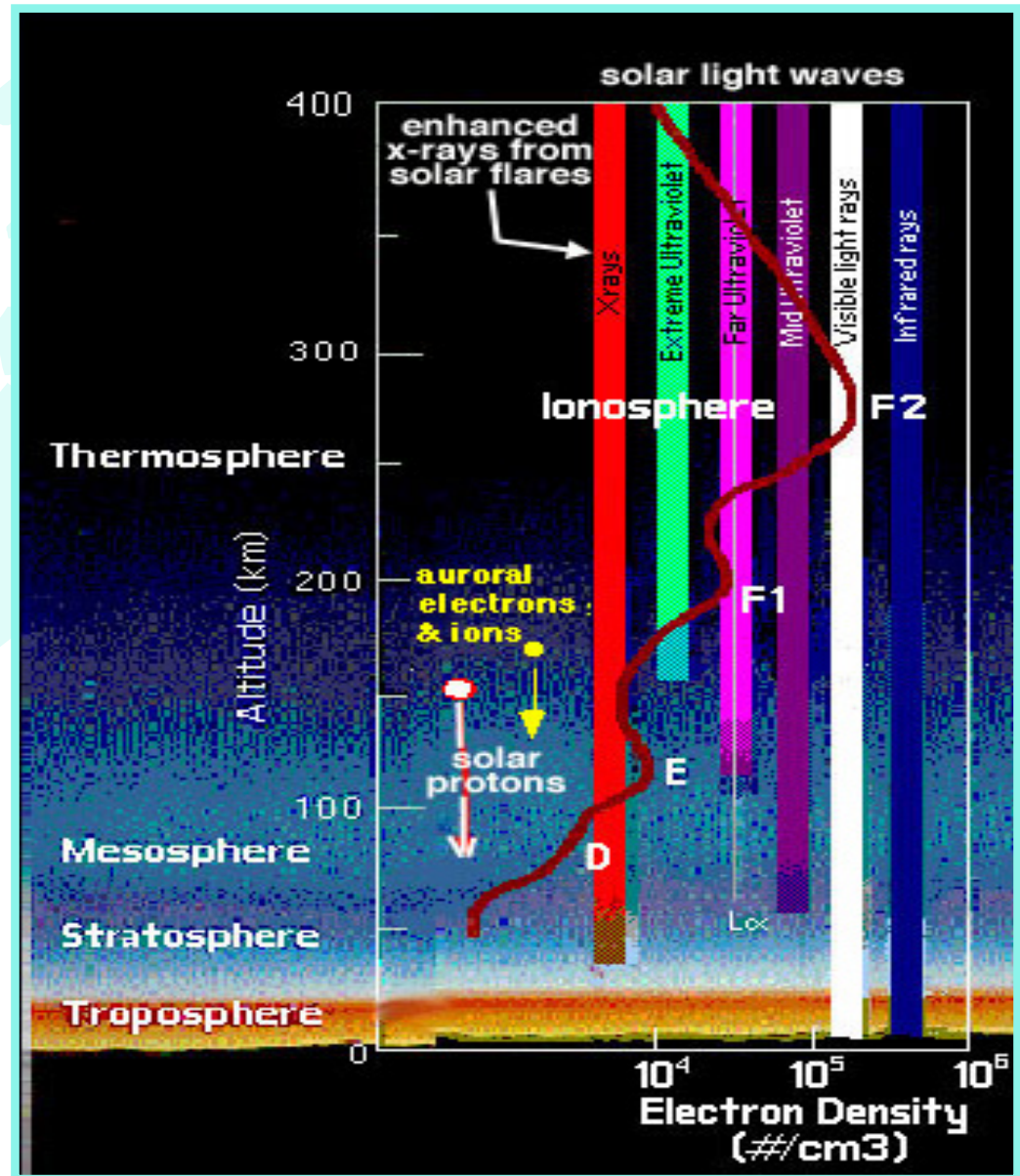
Joule heating and precipitating electrons coupled with Earth magnetosphere - important only at high latitudes and are a secondary global ionization source

Bremsstrahlung X-rays from braking electrons constrained by the Earth magnetosphere can penetrate to the mesosphere or below

Precipitating protons and α -particles from solar wind - penetrate to the mesosphere or below

Credit: Tobiska

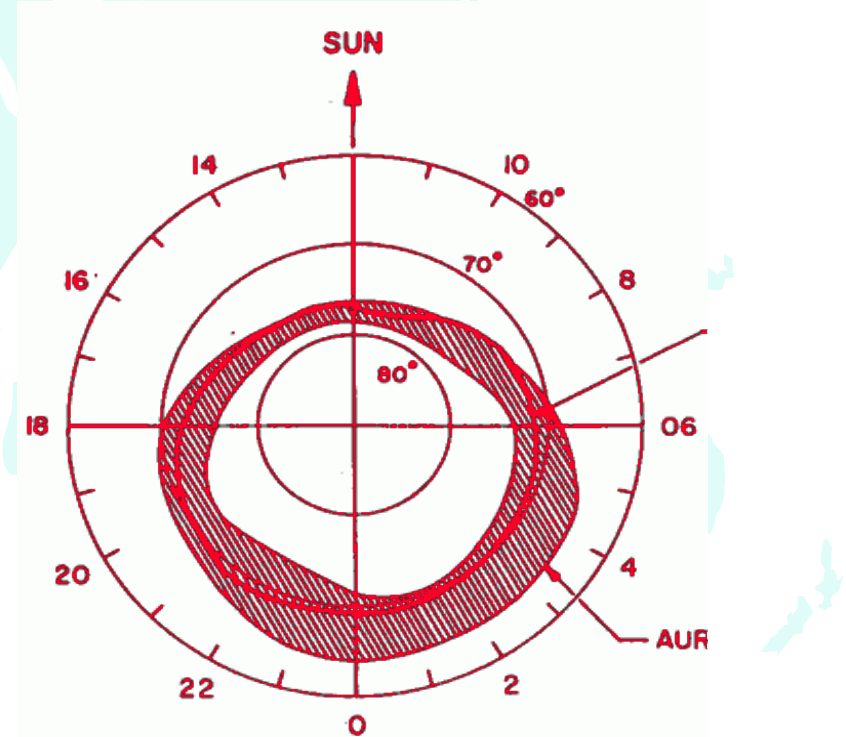
Solar radiation absorption compared to auroral and solar particles



Equation of Energy

Heat production: energetic charged particles are injected (precipitated) from the magnetosphere

Electrons spiral down 'open' magnetic field lines around auroral oval and impact neutral oxygen and nitrogen



Equation of Energy

Heat production: Joule Heating and Lorentz Forcing

- Ions accelerated by crossed **E** and **B** fields (left hand rule)
- **Joule heating**: macroscopic motions of ions and neutrals do not coincide → frictional heating
- **Lorentz Forcing**: microscopic motions of ions and neutrals do not coincide → collisional momentum transfer

Joule Heating

The macroscopic frictional heating of the thermosphere due to a bulk motion difference between the thermosphere and the ionosphere is Joule heating where E' is the perturbed electric field, V_N is the particle velocity (normal or perpendicular direction) B is the magnetic field, J is the electric current, and $J \cdot E$ is the ion heating

$$\mathbf{J} \cdot \mathbf{E}' = \mathbf{J} \cdot \left(\mathbf{E} + (\mathbf{V}_N \times \mathbf{B}) \right) \quad (6-6)$$

Lorentz forcing

Microscopic collisional momentum (hence energy) is transferred between **high speed ions and low speed neutrals** (or ‘ion drag’, when momentum is transferred from **high speed neutrals to low speed ions**)

$$-\rho v_{ni} (\mathbf{V}_n - \mathbf{V}_i) = \mathbf{V}_n \cdot (\mathbf{J} \times \mathbf{B}) \quad (6-7)$$

Joule vs Lorentz

At high altitudes
and/or large
electric fields
Joule heating is
dominant over
Lorentz by
approximately
an order of
magnitude

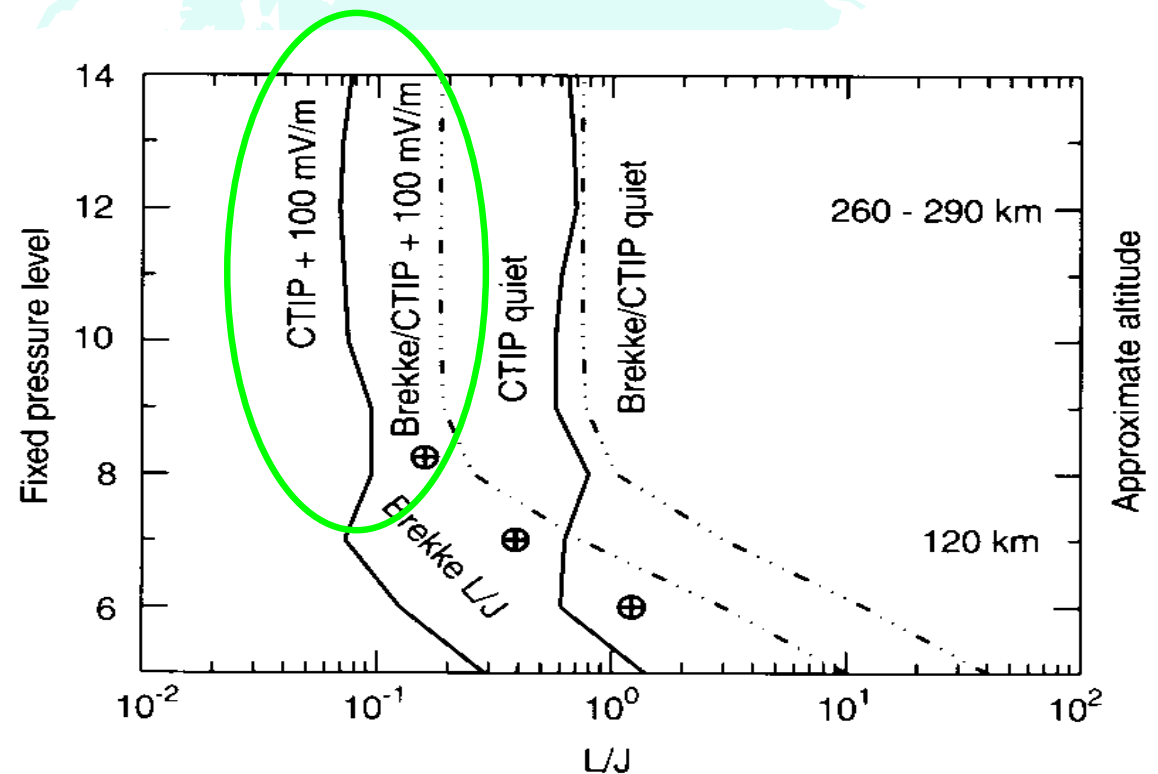
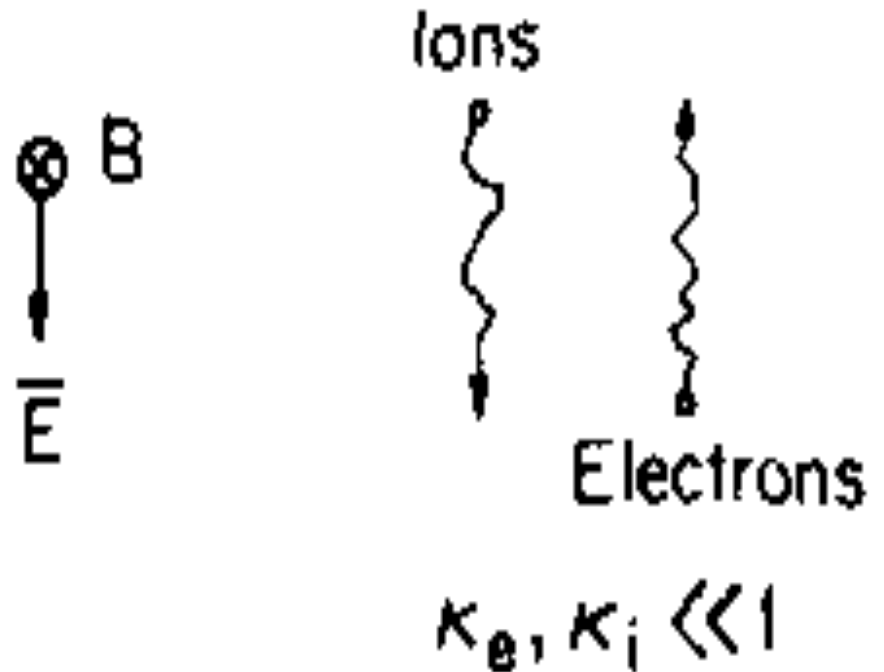


Fig. 3. Plot of the ratio of Lorentz forcing to Joule heating against height at latitude 70°N, longitude 162°E, Day 1. The *two solid lines* show the coupled model results for (a) quiet atmosphere, and (b) during the 100 mV/m enhanced electric field. The *broken lines* show results from the extrapolation of the Brekke formula, Eq. (4), for both quiet and enhanced atmospheres, whilst the *three discrete points* are the Brekke results

Why do ions drag?

In a collision-dominated atmosphere, e.g., below ~ 120 km, the trajectory is continually being re-established along the potential (gravity) gradient

Collisional Case



Why do ions drag?

(b) Collisionless Case

Ions



Electrons



$$\kappa_e, \kappa_i \gg 1$$

But in a collisionless atmosphere, above ~ 150 km, they are forced to spiral around electric field lines (left hand or motor rule)

Why do ions drag?

And, of course, there is a transition region - in the lower thermosphere near the 120-150 km altitude

(c) Intermediate Case



Equation of Energy

Heat production: winds (advection)

- Advection - non-cyclical motion of gas particles (as opposed to convection)
- Neutral winds transport energy
- Chemical transport by neutral winds allows species to be created by dissociation, for example, and be moved elsewhere before recombination

Equation of Energy

Heat production: wave dissipation

- Waves from a variety of sources will dissipate through turbulence, viscosity, resonance
- Possible sources of wave dissipation are
 - Thermospheric tides
 - Gravity waves aka atmospheric gravity waves
aka Travelling Atmospheric Disturbances
 - MHD waves

Equation of Energy

Heat production: molecular conduction

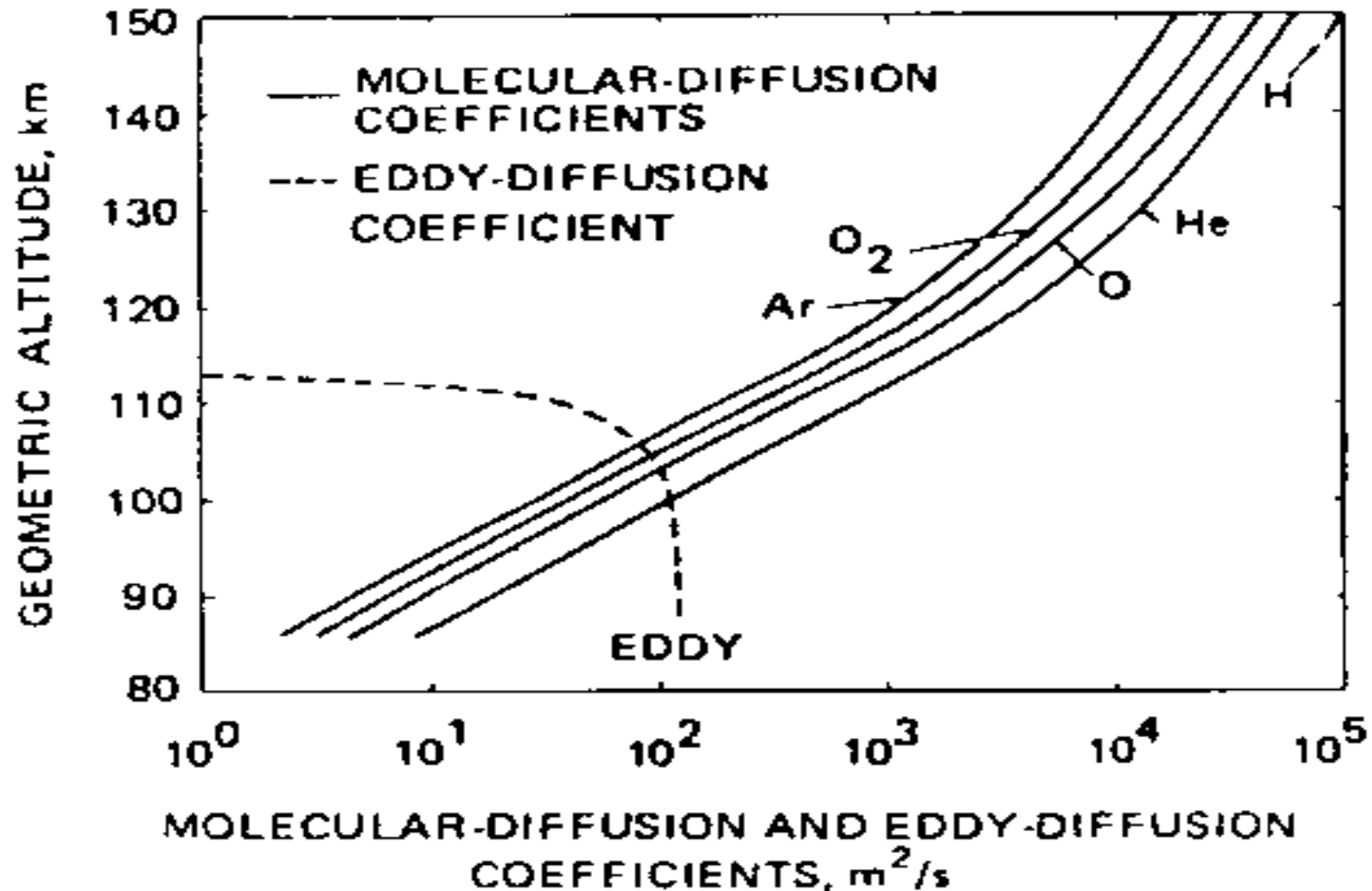
- Thermospheric temperature becomes asymptotic $> 400\text{-}500$ km due to rapid thermal conductivity
- Heat flux $\propto$ temperature gradient, i.e., $\Phi \propto \nabla T$

Equation of Energy

Heat production: eddy transport

- This is an important process below the turbopause (~ 120 km) and dominates below ~ 100 km
- Eddy mixing transports heat down to into the mesosphere (around 50 km)

Diffusion Coefficients



Vandenberg launch Sep 2005



Vandenberg launch



Vandenberg launch



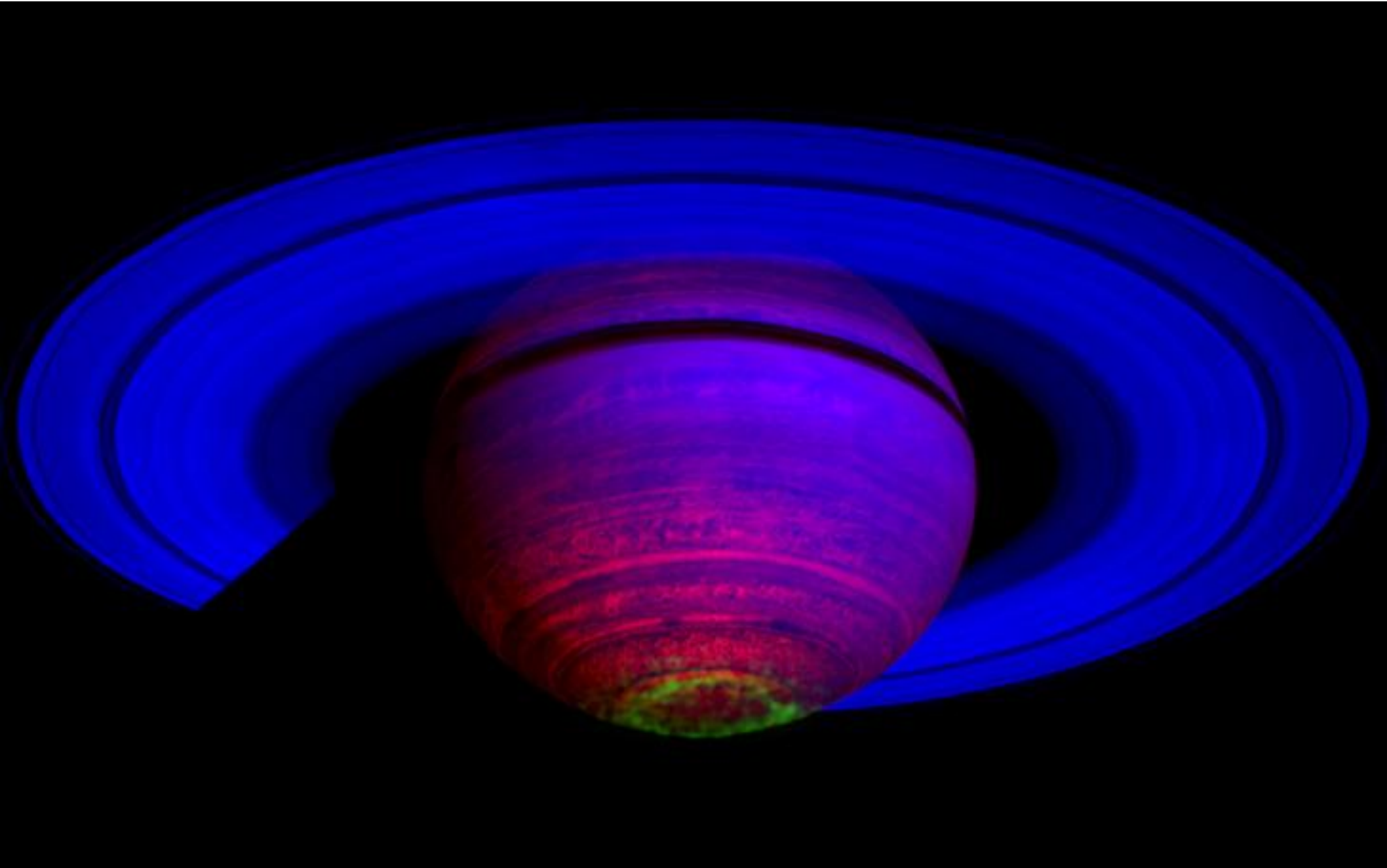
Equation of Energy

Heat loss: radiation

- This is the principal mechanism of thermospheric heat loss (CO_2 and NO)
- Auroral emission (spectral, e.g. OH lines and Oxygen line @ 630 nm red line and 557 nm green line)
- Airglow emission (e.g. O, N)







The Continuity Equation

Ion Production

Ions and electrons, once formed, will tend to recombine unless they are prevented from doing so and we can rewrite the continuity equation so that



$$\frac{\partial n}{\partial t} = P - L - \nabla \cdot (n\mathbf{V}) \quad (6-8)$$

Ion Production

Ion Transport

$$\frac{\partial n}{\partial t} = P - L - \nabla \cdot (n \mathbf{V}) \quad (6-9)$$

Ions and electrons will be influenced by the vertical pressure gradient and the heavier ions will tend to 'settle'

The resultant separation will form an electric field which will tend to restore the particles

Thermospheric chemistry

The dominant neutral species are atomic oxygen, molecular nitrogen, molecular oxygen, helium, and hydrogen, respectively

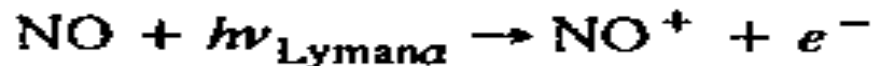
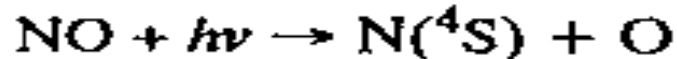
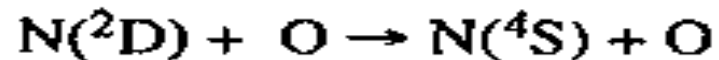
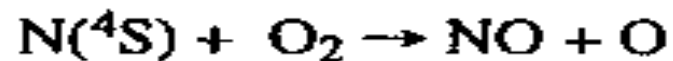
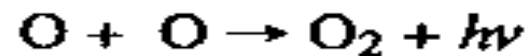
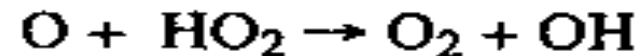
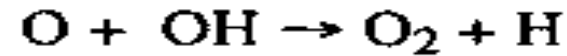
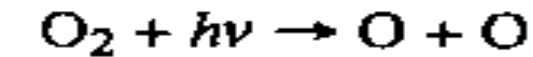
The dominant ion species is atomic oxygen

Much of the chemistry involves O and N in one form or another although nitrogen has a very small photodissociation cross-section

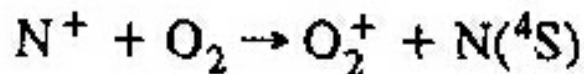
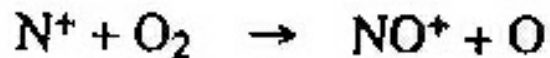
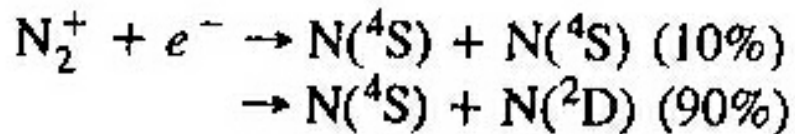
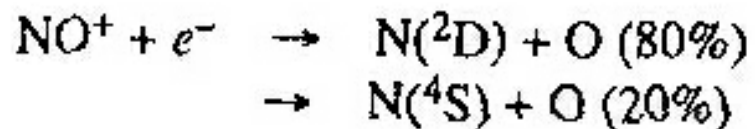
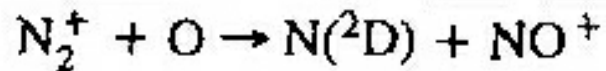
Ionization potentials

Atom or molecule	H	He	C	N	O	Ar	Ne	NO	O ₂	CO ₂	N ₂
Ionization Potential (eV)	13.60	24.59	11.26	14.53	13.62	15.76	21.56	9.25	12.08	13.79	15.58
Max wavelength (nm)	91.2	50.4	110.1	85.3	91.0	78.7	57.5	134.0	102.7	89.9	79.6

Neutral-Neutral Chemistry (photodissociative reactions)

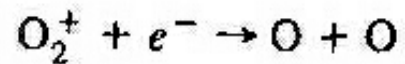


Neutral-Ion Chemistry

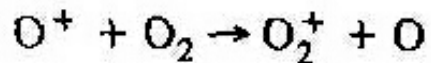
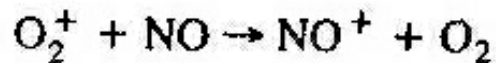
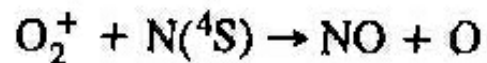


Dissociative recombination
(fast reactions)

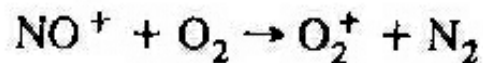
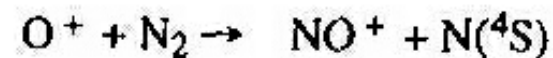
Neutral-Ion Chemistry



Dissociative recombination
(fast reactions)



Charge Exchange

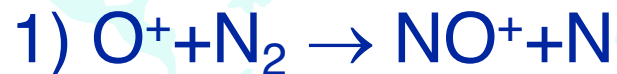


$$T_i = 0.667 T_i + 0.333 T_n; T_i \text{ is ion temperature}$$

and $T_R = (T_i + T_n)/2$.

F2 Region Chemistry

For the dominant ionospheric layer, the F2 region, there is a two-stage recombination process



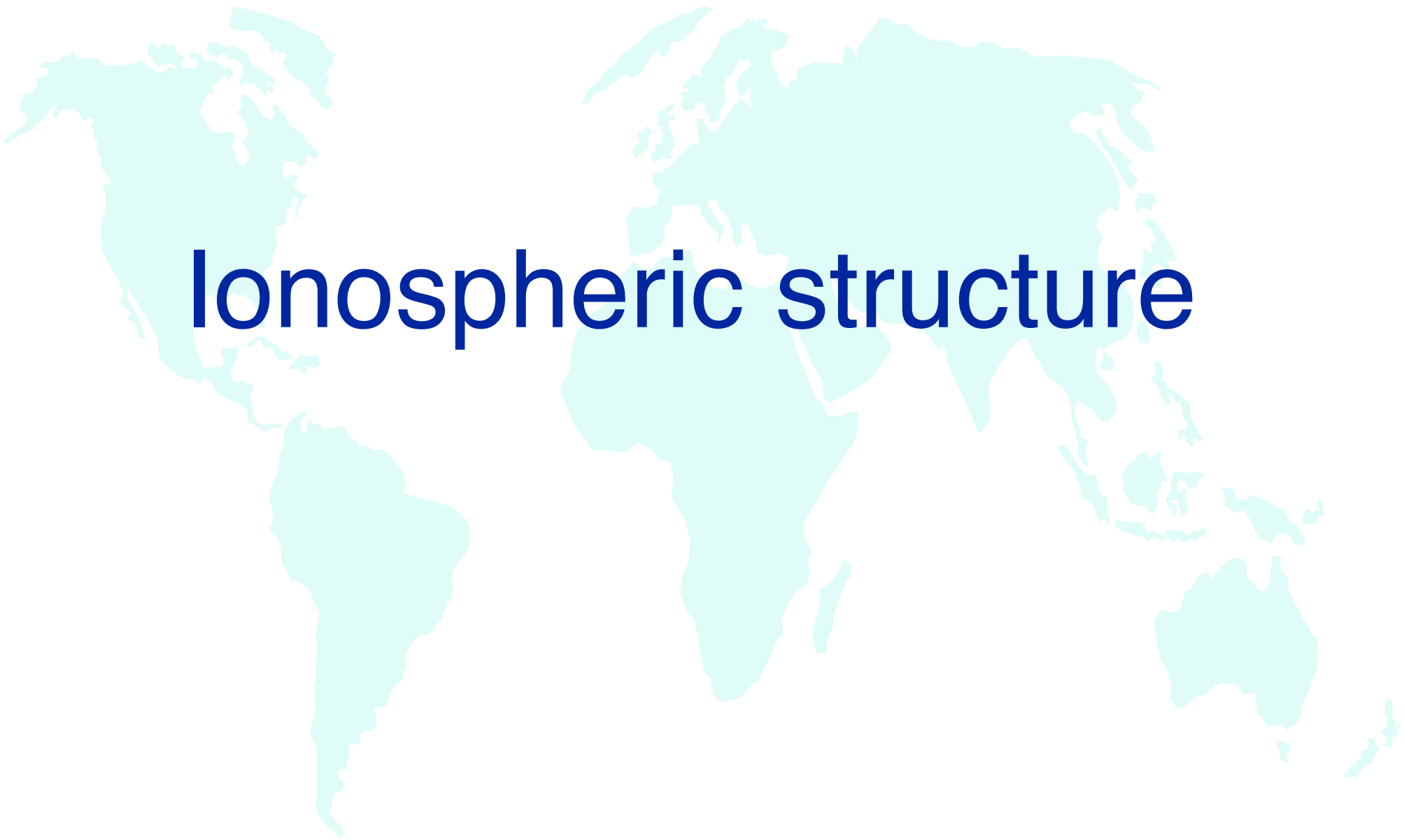
where the rate is $\propto \beta [\text{O}^+]$ which controls the rate at high altitudes (reaction between most predominant species)



where the rate is $\propto \alpha [\text{NO}^+][n_e]$ which controls the rate at low altitudes where α is large (more NO available)

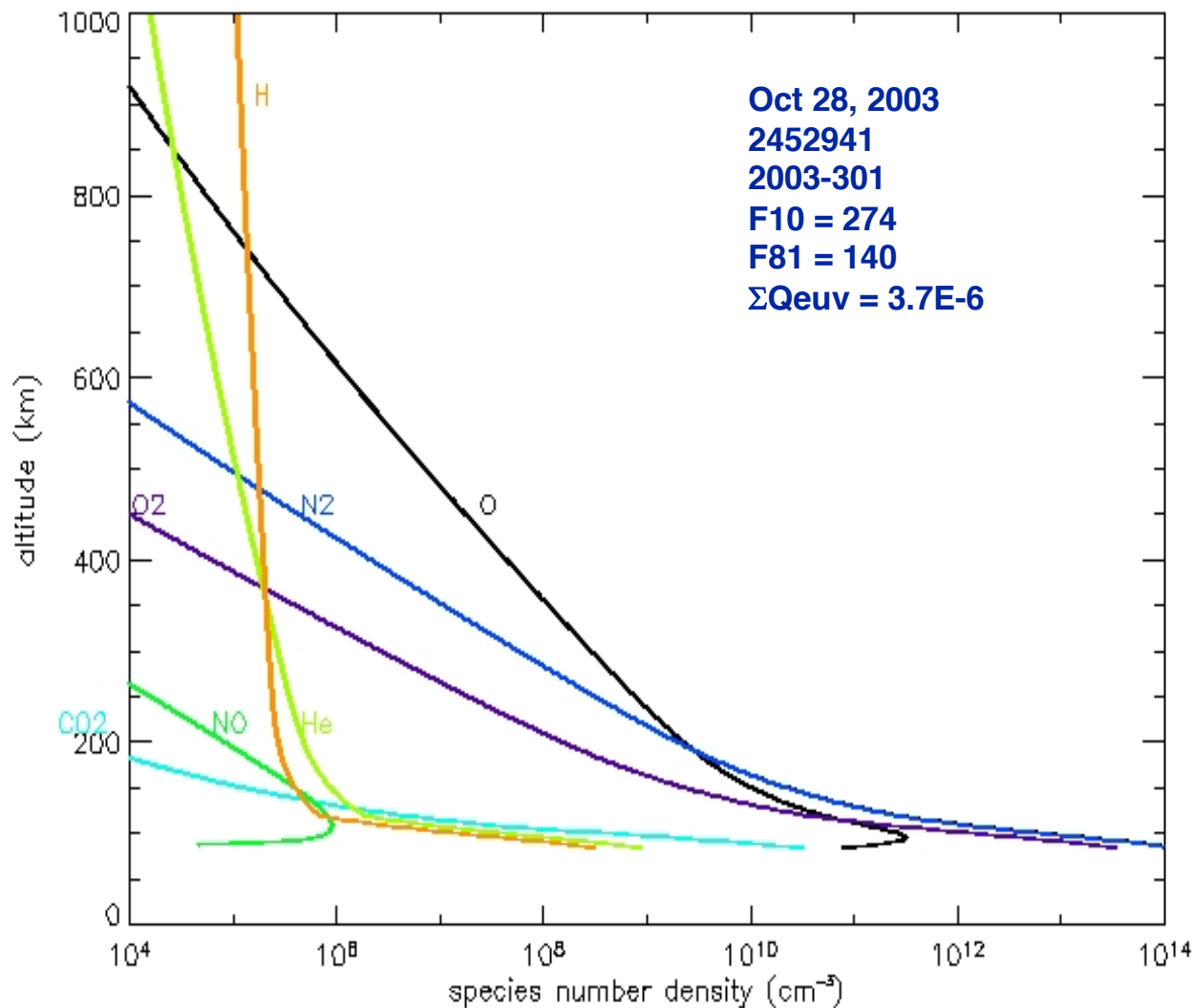
A light blue silhouette of a world map is centered in the background of the slide.

BREAK

A light blue silhouette of a world map serves as a background for the title text.

Ionospheric structure

Thermosphere species



Major species

O, O₂, N₂

Minor species

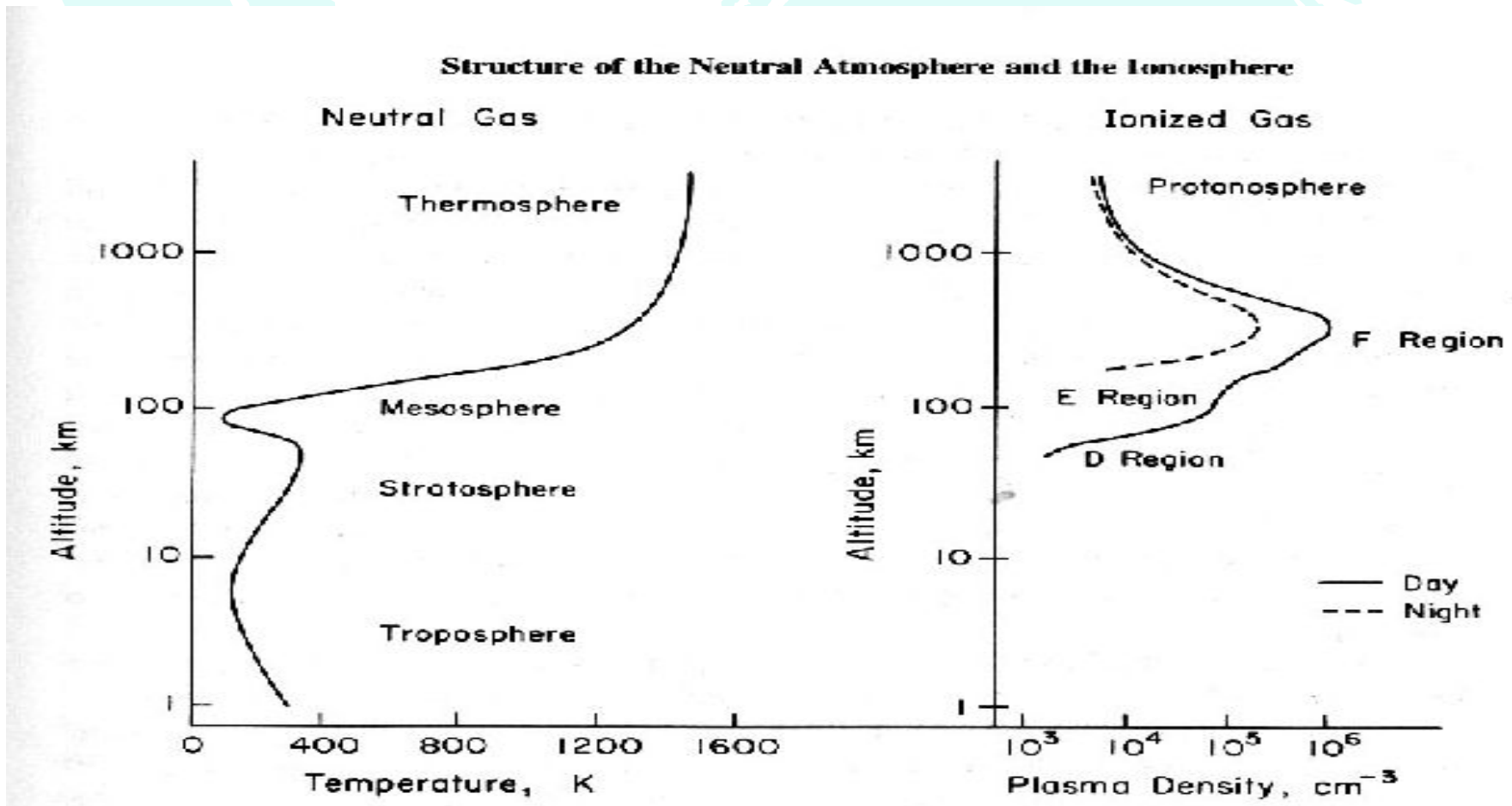
NO, CO₂ are significant for cooling

He, H useful for aeronautical applications

Ar

Credit: Tobiska

Thermosphere-Ionosphere structure



Thermosphere vs Ionosphere

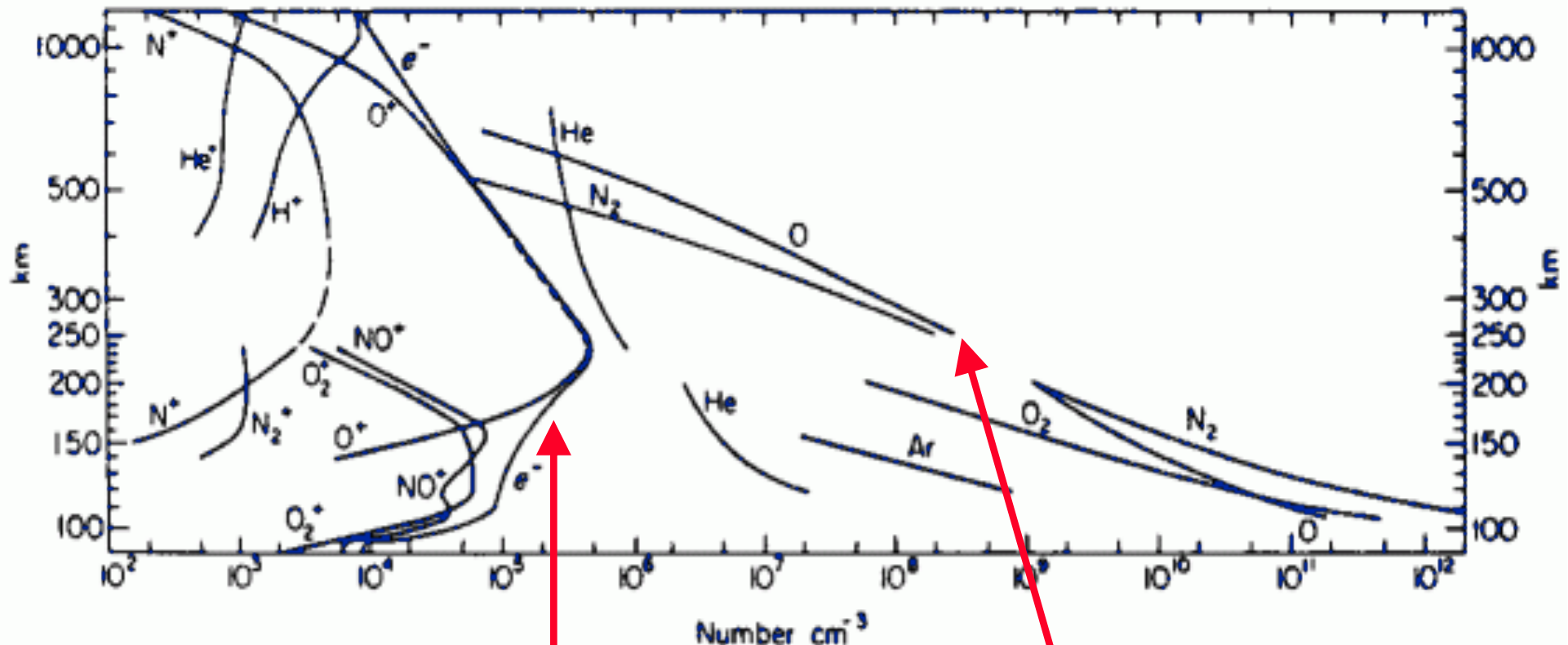


Fig. 1.2. International Quiet Solar Year (IQSY) daytime atmospheric composition, based on

Ionospheric density $\ll$ Thermospheric density by $\sim 1\text{E}5$

Ionospheric Layers

- ◆ D-region:
 - 50-85 km
 - dominated by H_2O^+
 - $\lambda < 1 \text{ nm}$ (Xray) or Lyman-alpha (121.6 nm)
- ◆ E-region:
 - 85-140 km
 - dominated by O_2^+
 - $1 < \lambda < 10 \text{ nm}$ (XUV) or Lyman-beta (102.6 nm or CIII (97.7 nm)
 - E-F1 valley is the cross-over between the two regions

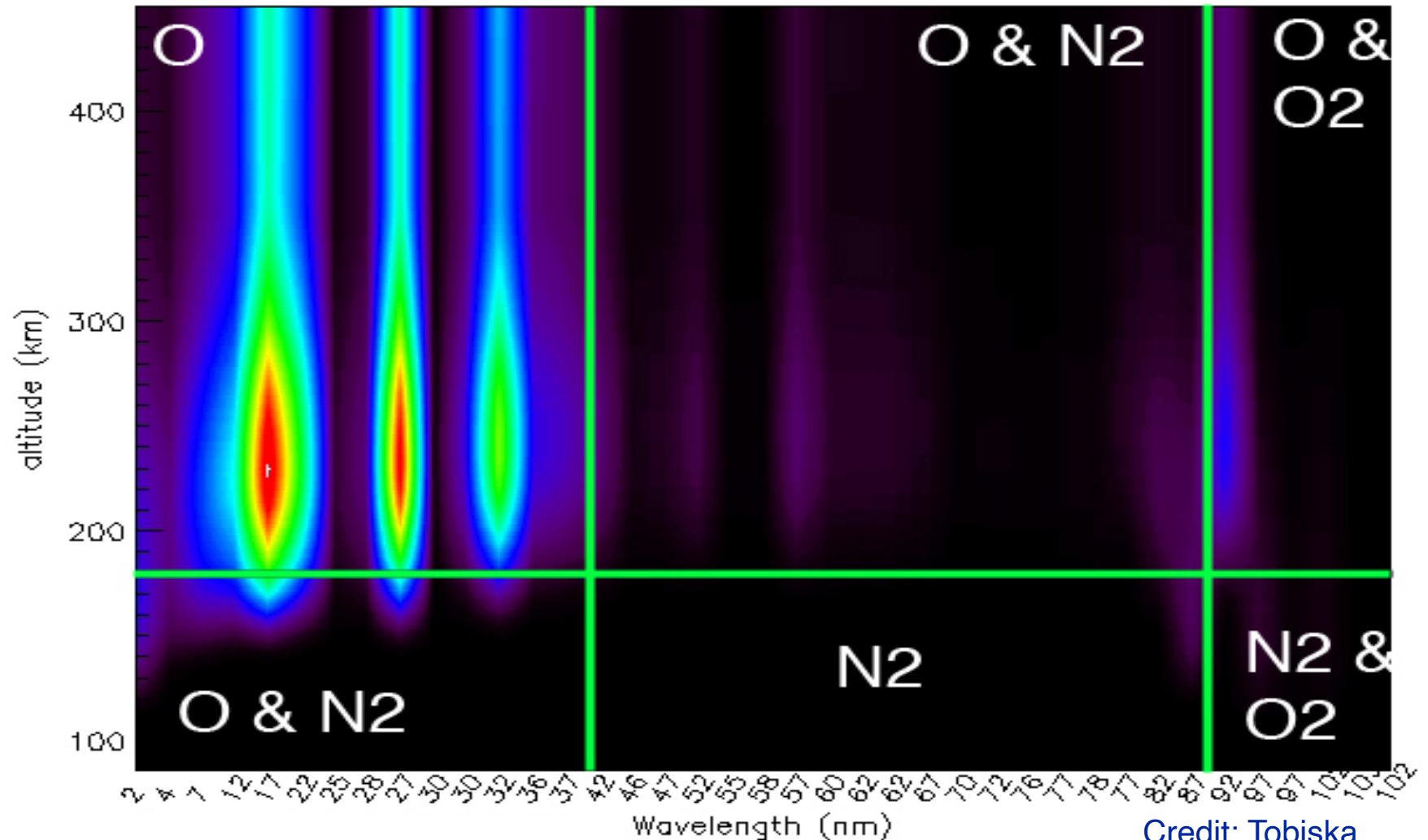
Ionospheric Layers

- ◆ F1-region:
 - 140-250 km
 - dominated by NO^+
 - $10 < \lambda < 91 \text{ nm}$ & 121 nm (EUV and Lyman-alpha)
 - F1 maximum electron density ~200 km
- ◆ F2-region:
 - 250-500 km
 - dominated by O^+
 - $10 < \lambda < 91 \text{ nm}$ (EUV)
 - F2 maximum electron density ~300 km (F2max)

Ionospheric Layers

- ◆ Topside ionosphere:
 - >500 km
 - dominated by O^+
 - $\lambda < 102.6$ nm (EUV < Lyman-beta)
- ◆ Total Electron Content (TEC):
 - integrated ionosphere from top to bottom

Photoionization regions

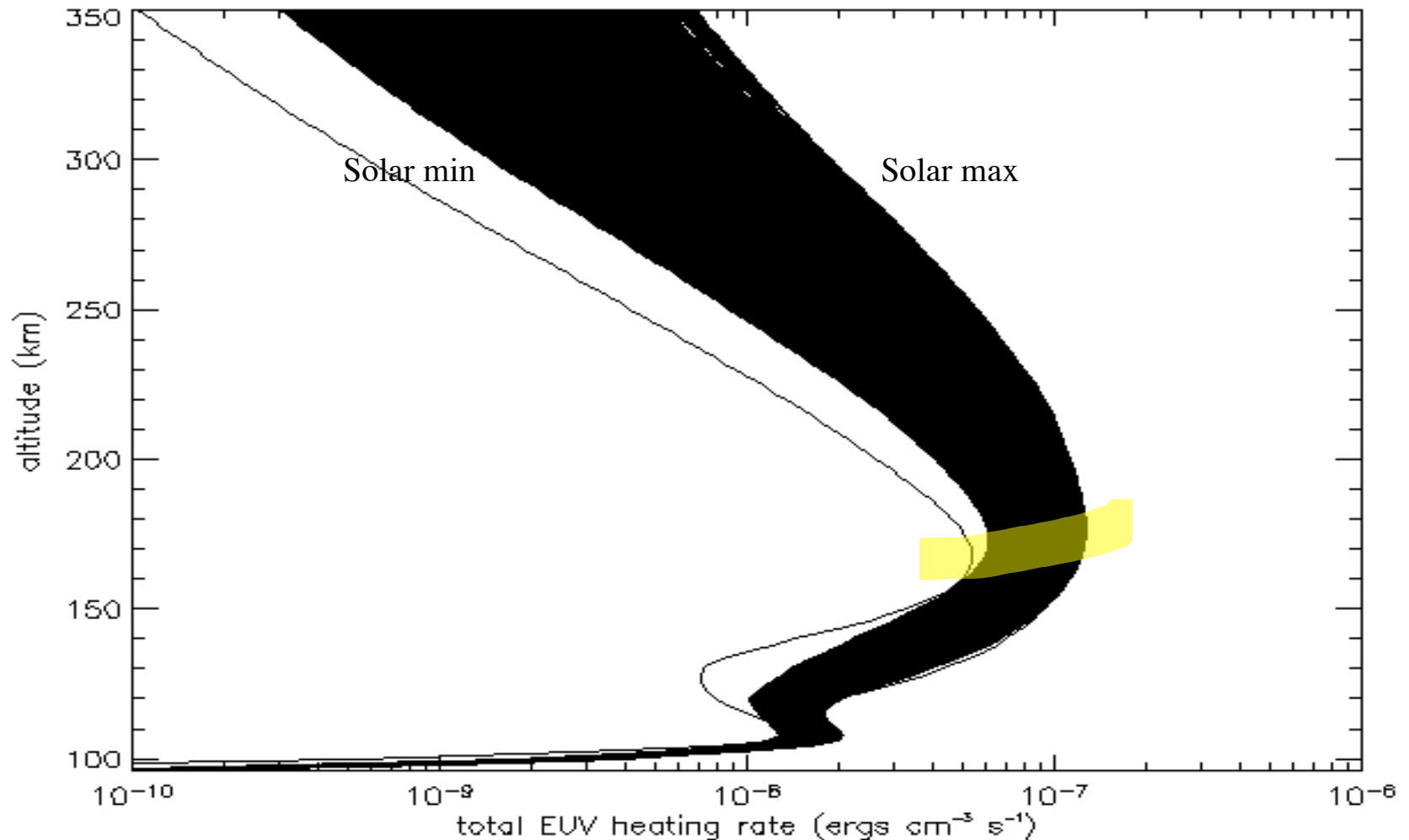


Credit: Tobiska

Ionospheric Layers

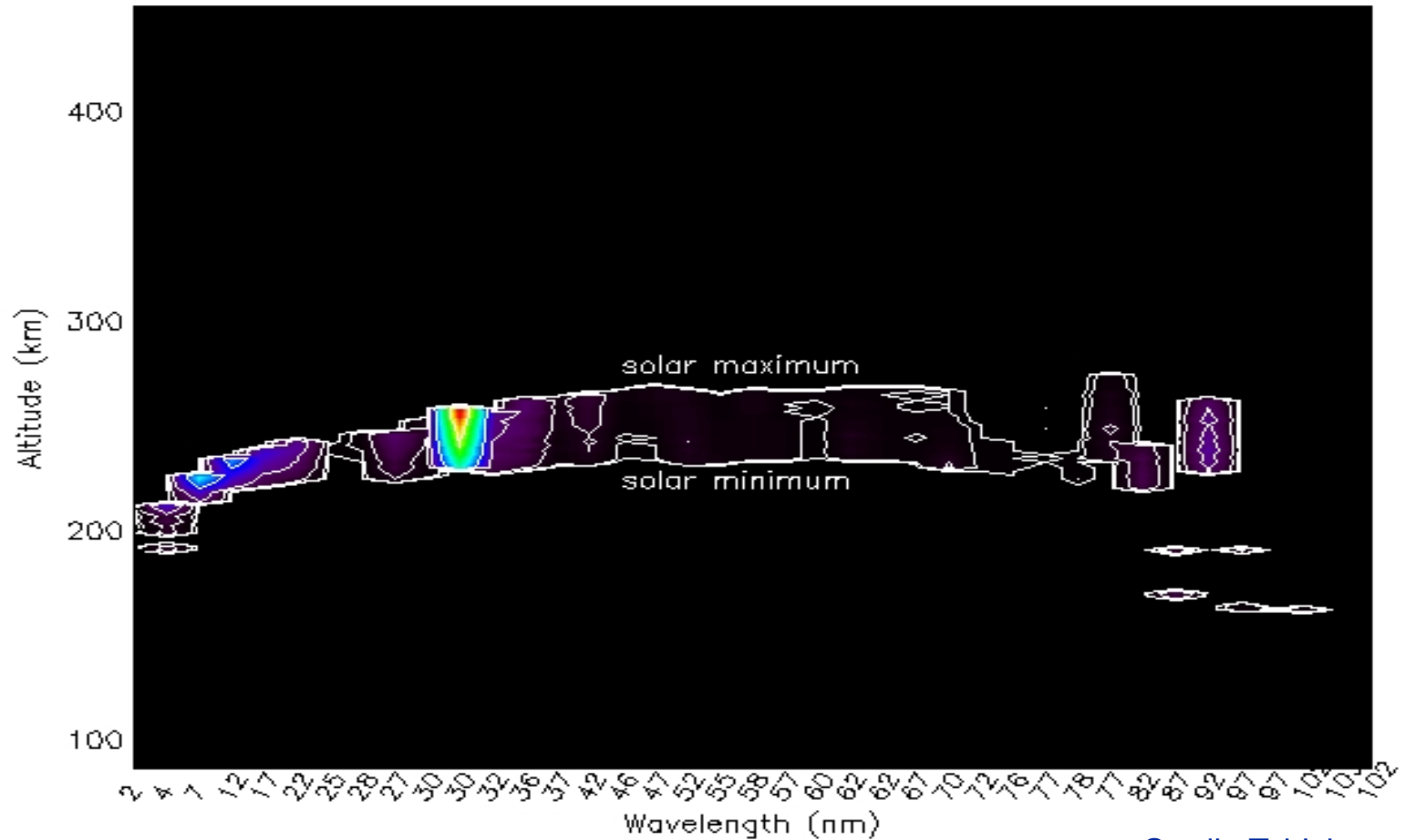
- ◆ Additional concepts:
 - Relatively little N_2^+ because fast recombinations tend to form NO^+
 - Hotter thermosphere (during higher solar activity) tends to raise peak layers of ionization because there is a denser atmosphere at a given altitude and the unit optical depth is higher.

Solar max-min heating



Credit: Tobiska

Solar max-min optical depth



Credit: Tobiska

Scale Height

In equilibrium and considering only gravity and the pressure gradient

$$dp = -nmg \, dh \quad (6-13)$$

$$\frac{dp}{dh} = -nmg \quad (6-14)$$

but $p = nkT$

$$\frac{1}{p} \frac{dp}{dh} = -\frac{mg}{kT} = -\frac{1}{H} \quad (6-15)$$

where $H = kT/mg$ is the scale height

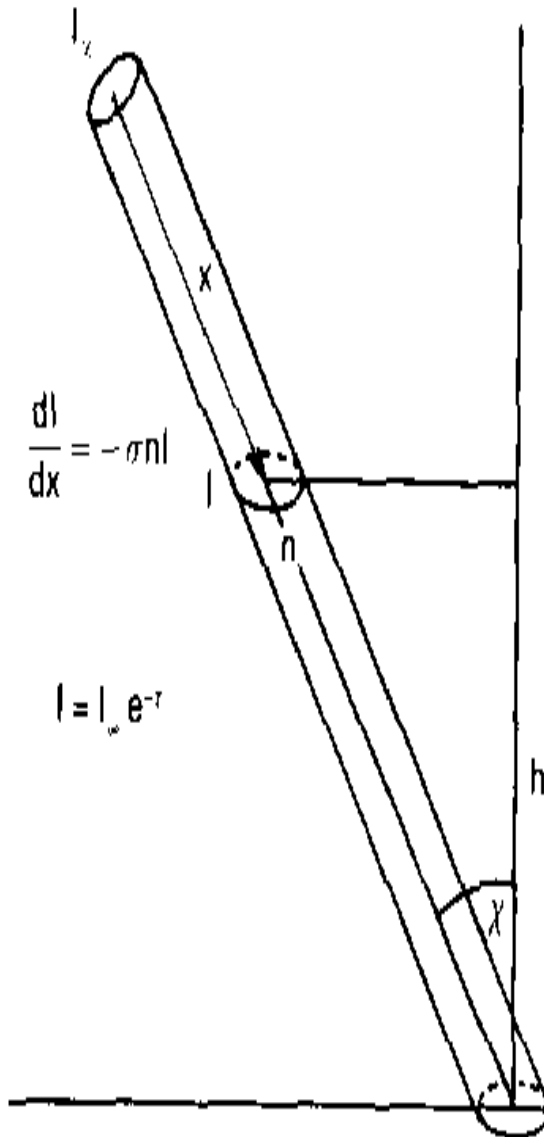
Chapman Layer Production

Assume single species gas with exponential $N(h)$, constant H , plane stratified, and monochromatic radiation for ion production rate, P

$$P = P_{m_0} \exp(1 - z - \exp(-z) \sec \chi) \quad (6-17)$$

$$\text{where } z = \frac{h - h_0}{H}, \chi = \text{SZA},$$

m_0 denotes maximum ion production



Chapman Function

Pedersen (1927) and Chapman (1931) provided a mathematical form for relating ion production with ionizing radiation

$$P = \eta \sigma n I \quad (6-16)$$

P = Production rate, I = intensity of ionizing radiation, n = number density (concentration) of neutral particles, σ = absorption cross-section, η = ionization efficiency

Chapman Layer Loss

$$L = \alpha [X^+] n = \alpha n^2 \quad (6-18)$$

Loss rate

Ion density

Recombination coefficient

for ion X at a number density (concentration), n , with a recombination coefficient, α

At equilibrium, Production = Loss

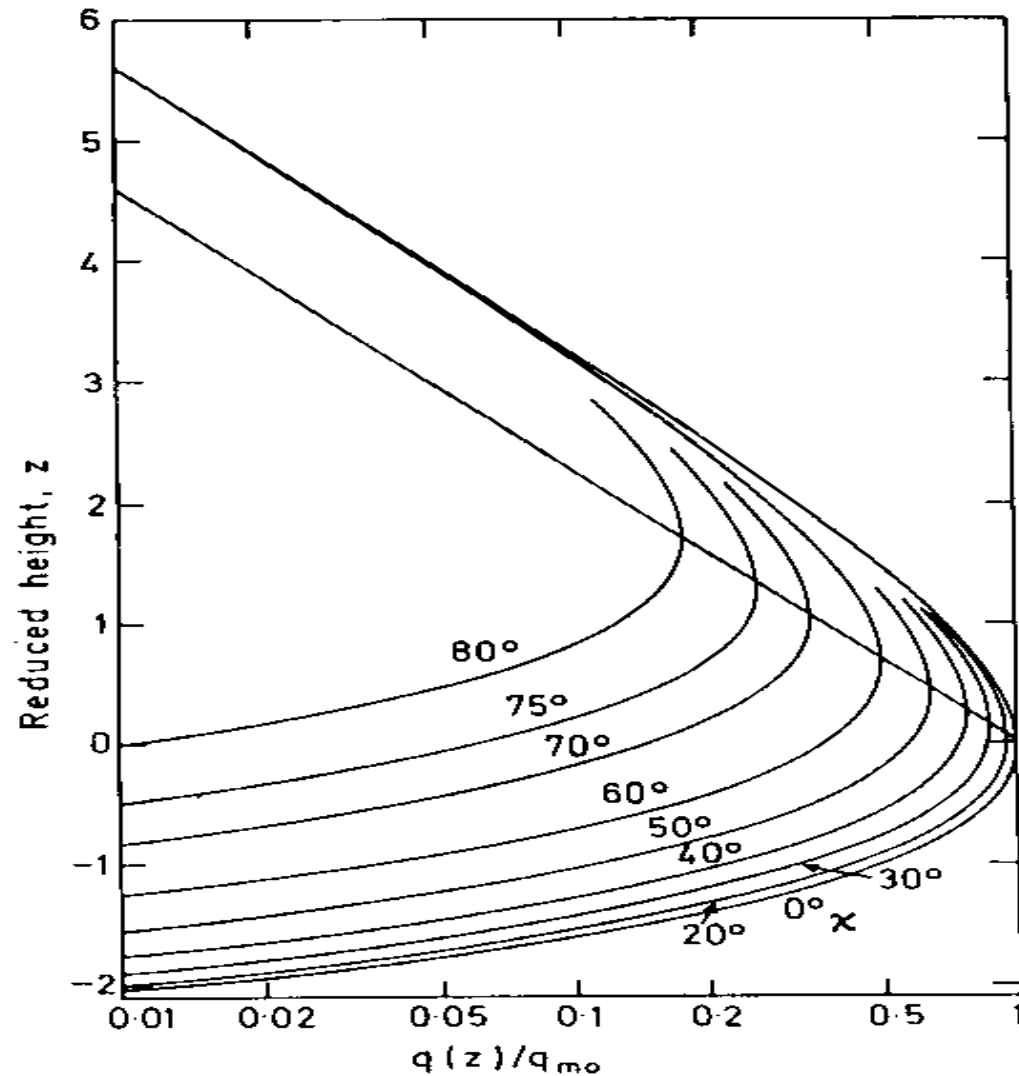
Chapman Layers

If $P \propto I_n$ and as h (altitude) increases, ionizing photons, I , increase and concentration, n , decreases, we expect that *at some particular height*, there is a maximum production rate and hence a maximum ion density.

Chapman Profile

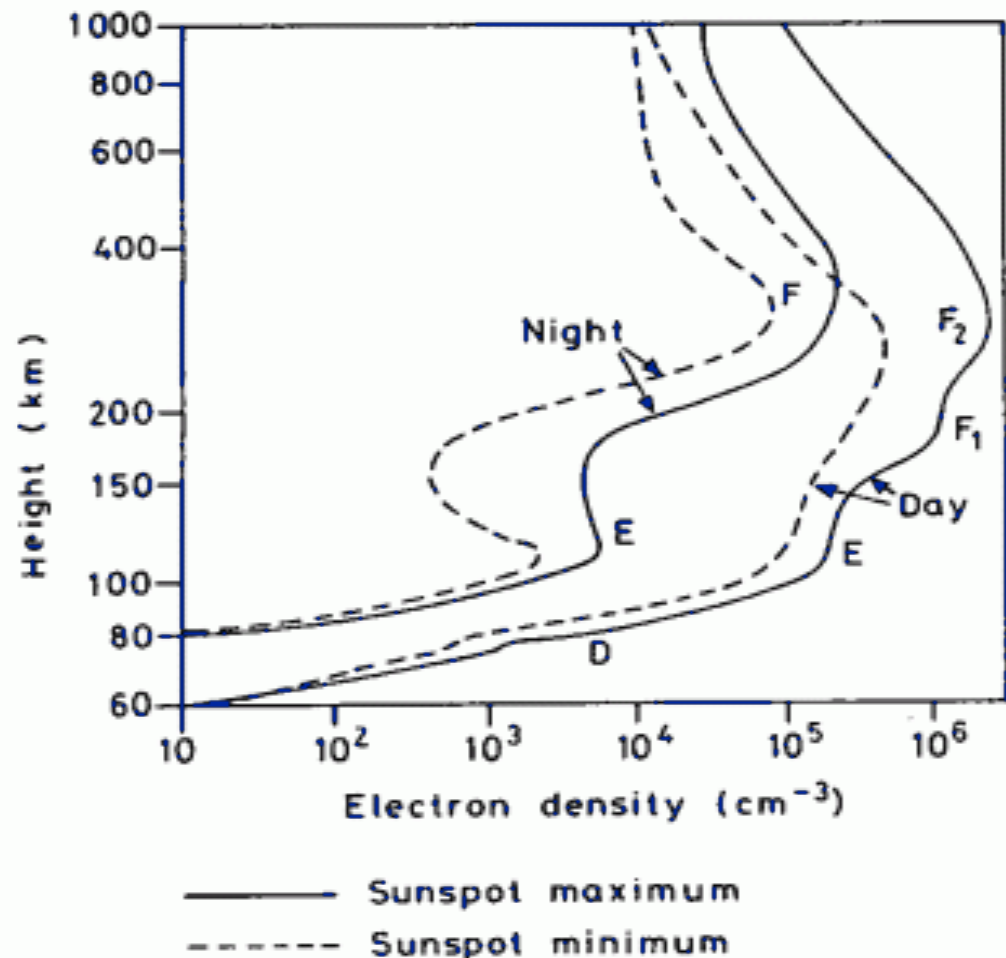
($q = P$ in this figure)

X is the angle of incidence of the solar photons where 90 degrees is the limb and 0 degrees is overhead



Ionosphere structure

- ◆ Exists inside (and above) the thermosphere, made up of free electrons and positive ions (and a few negative ions)
- ◆ Ionosphere is electrically **NEUTRAL** overall (but not necessarily neutral locally)



Ionosphere structure

- ◆ Atomic Ions: O^+ , N^+ , H^+ , He^+
- ◆ Molecular Ions: O_2^+ , N_2^+ , NO^+

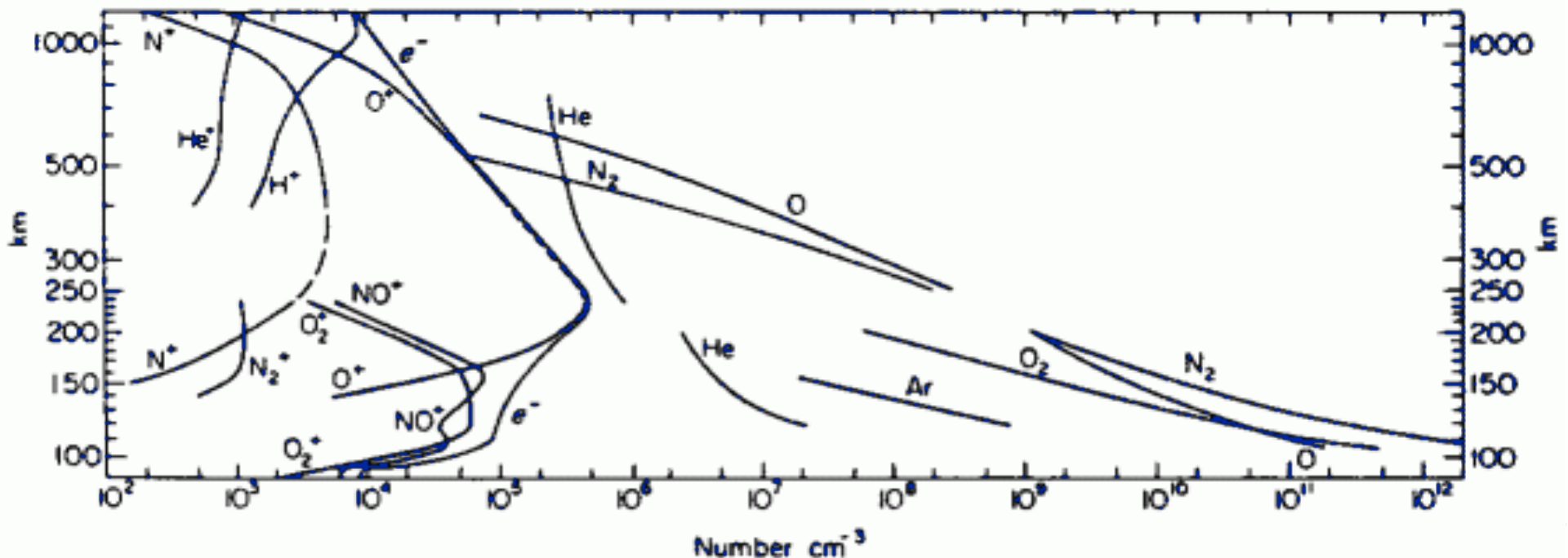


Fig. 1.2. International Quiet Solar Year (IQSY) daytime atmospheric composition, based on

Chapman Layers

The production rate is greatest at the level where the optical depth is unity.

We expect an ‘ α -Chapman layer’ at ~ 180 km – the F1 layer.

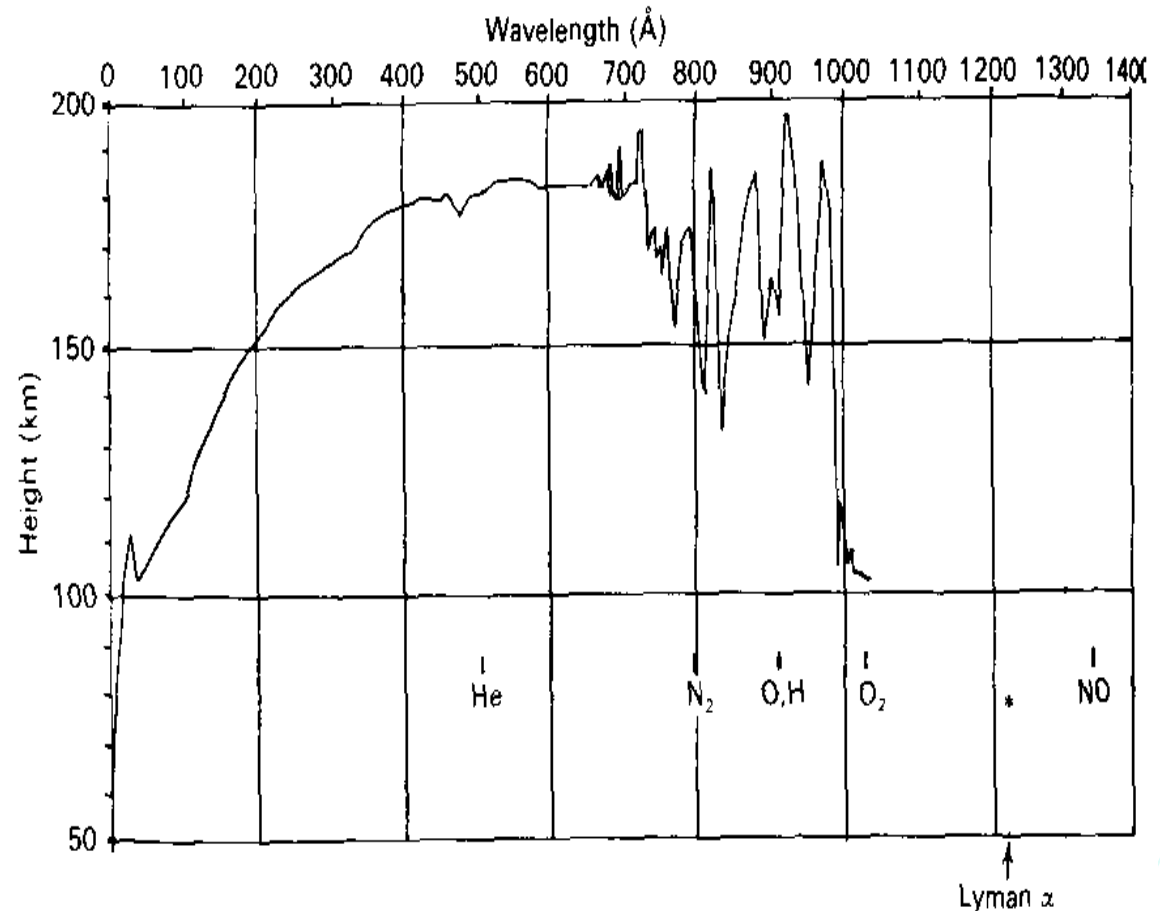


Fig. 6.11 Height where the optical depth reaches unity for radiation vertically incident on a

Complications with the ionosphere loss processes

Electrons may combine with neutrals to form negative ions

Positive ions may also combine with negative ions as well as with electrons

$$N_m = N_{m0} \cos \chi \quad (6-19)$$

This is a β -Chapman layer (modified recombination coefficient β) for N_m = to the maximum concentration of ions

Equation of Motion

- ◆ The lower thermosphere is collision-dominated
- ◆ Basic equations of fluid dynamics may be used, e.g. Navier-Stokes
- ◆ We resort to Newtonian mechanics,

$$F = ma$$
to describe the forces acting upon a *packet of neutral gas* and its rate of change of velocity

Equation of Motion

Take a single packet of gas and draw a force diagram in the Lagrangian frame i.e., moving with the gas packet

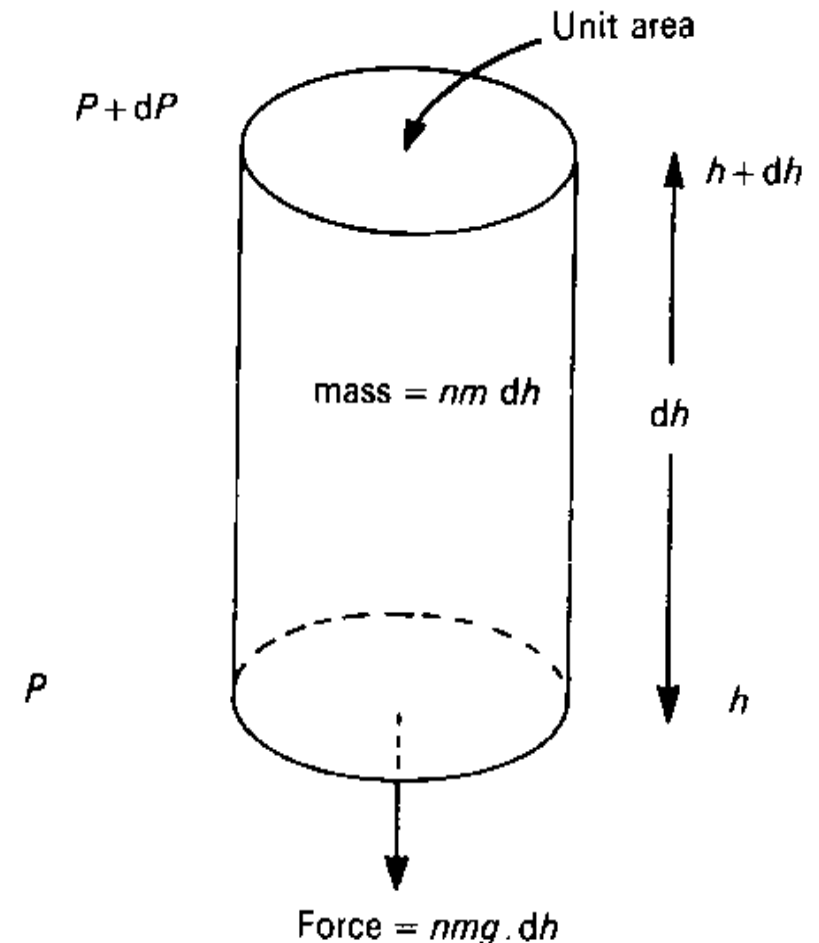
Acting upon the gas packet:

Gravity & Pressure Gradient

Coriolis effects

Ion Drag

Viscosity



Equation of Motion

acceleration = force ÷ unit mass

$$a = dU/dt = F/m$$

$$\frac{dU}{dt} = g - \underbrace{\frac{1}{\rho} \nabla p}_{\text{pressure gradient}} - \underbrace{2\Omega \wedge U}_{\text{Coriolis forcing}} - \underbrace{v_{ni}(U - V)}_{\text{ion drag}} + \underbrace{\frac{1}{\rho} \nabla(\mu \nabla U)}_{\text{viscosity}} \quad (6-10)$$

U , ρ , p , v_{ni} = neutral gas velocity, mass density, pressure, neutral-ion collision frequency; Ω = Earth's angular rotation; V = ion velocity; μ = viscosity coefficient, g = gravity

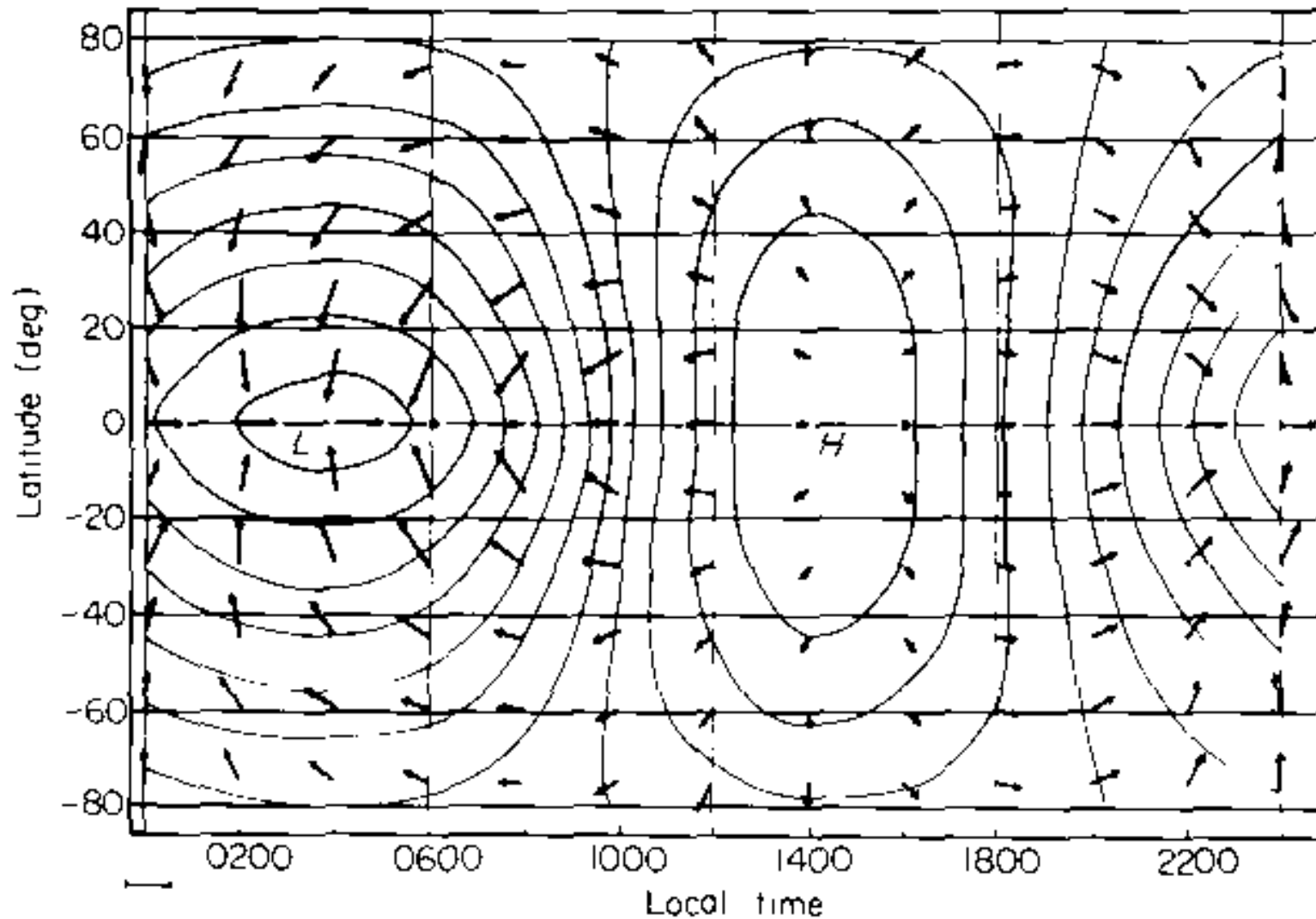
How important are each of these terms?
Depends on location in thermosphere.

Equation of Motion

- ◆ Convert to Eulerian frame, i.e. in the frame of the Earth, by means of a canonical transform

$$\underbrace{\frac{dX}{dt}}_{\substack{\text{rate of change of } X \\ \text{(a gas parameter)} \\ \text{in the packet of gas}}} = \underbrace{\frac{\partial X}{\partial t}}_{\substack{\text{rate of change of } X \\ \text{at a fixed point on} \\ \text{the Earth's surface}}} + \underbrace{(V \cdot \nabla) X}_{\substack{\text{advection} \\ \text{term}}} \quad (6-11)$$

Thermal Winds



Tides and Winds Terminology

In the upper thermosphere, solar heating of the thermosphere itself instigates horizontal winds flowing from the dayside to the nightside

The Earth's rotation implies that these are 'tidal' – however we refer to them as 'winds', since in the lower thermosphere ($< \sim 120$ km), vertical winds propagate upwards from the mesosphere

These are referred to as (thermospheric) 'tides'

Tides and Winds

The sun (and moon) will cause tidal effects that propagate westwards, due to heat and gravity

The solar forcing and natural modes (Hough modes) of oscillation (cf. guitar string) will combine to give tidal modes

Hough function (n,m): n cycles per day, m-n nodes between the poles

The dominant tidal modes (in order) are solar-diurnal, solar-semidiurnal (heating-drive), and lunar-diurnal (gravity-driven)

Thermal Winds

We would expect an eastward wind
across the sunset terminator

At the equator, Coriolis terms $\rightarrow 0$

To a first approximation, the neutral
atmosphere is slowed by viscosity and
we ignore ion drag

Below the turbopause, it is eddy viscosity

Above the turbopause, it is molecular viscosity

Thermospheric Waves

Waves propagate from any disturbance.

The restoring forces are both gravity vertically (e.g. ocean waves) and compression horizontally (e.g. sound wave): hence **‘acoustic gravity waves’** or **‘atmospheric gravity waves’** (AGWs), or more generally, **‘travelling atmospheric disturbances’** (TADs).

Thermosphere Waves

Wave propagation tends to be along
'wave guides'

An obvious wave guide is the pressure
gradient of the atmosphere (c.f. ocean
waves)

Phase propagation is upwards.

Thermosphere Waves

A disturbance will generally create a 'family' of TADs

Large Scale (LS)

Medium Scale (MS)

LS-TADs originate from magnetically conjugate auroral disturbances in the thermosphere

λ (wavelength) ~ 1000 s km, τ (period) ~ 10 s min, v (velocity) ~ 100 s m/s, particle motion $\leftrightarrow$

MS-TADs originate from the middle atmosphere

$\lambda \sim 10$ s km, $\tau \sim$ mins, $v \sim 10$ s m/s, wavefront $\sim 45^\circ$

Large Scale TADs

What is the triggering disturbance?

Large electric fields in the polar regions drive the ionosphere through $E \times B$ forcing (left hand or motor rule)

The ionosphere 'rubs against' the pseudo-stationary thermosphere

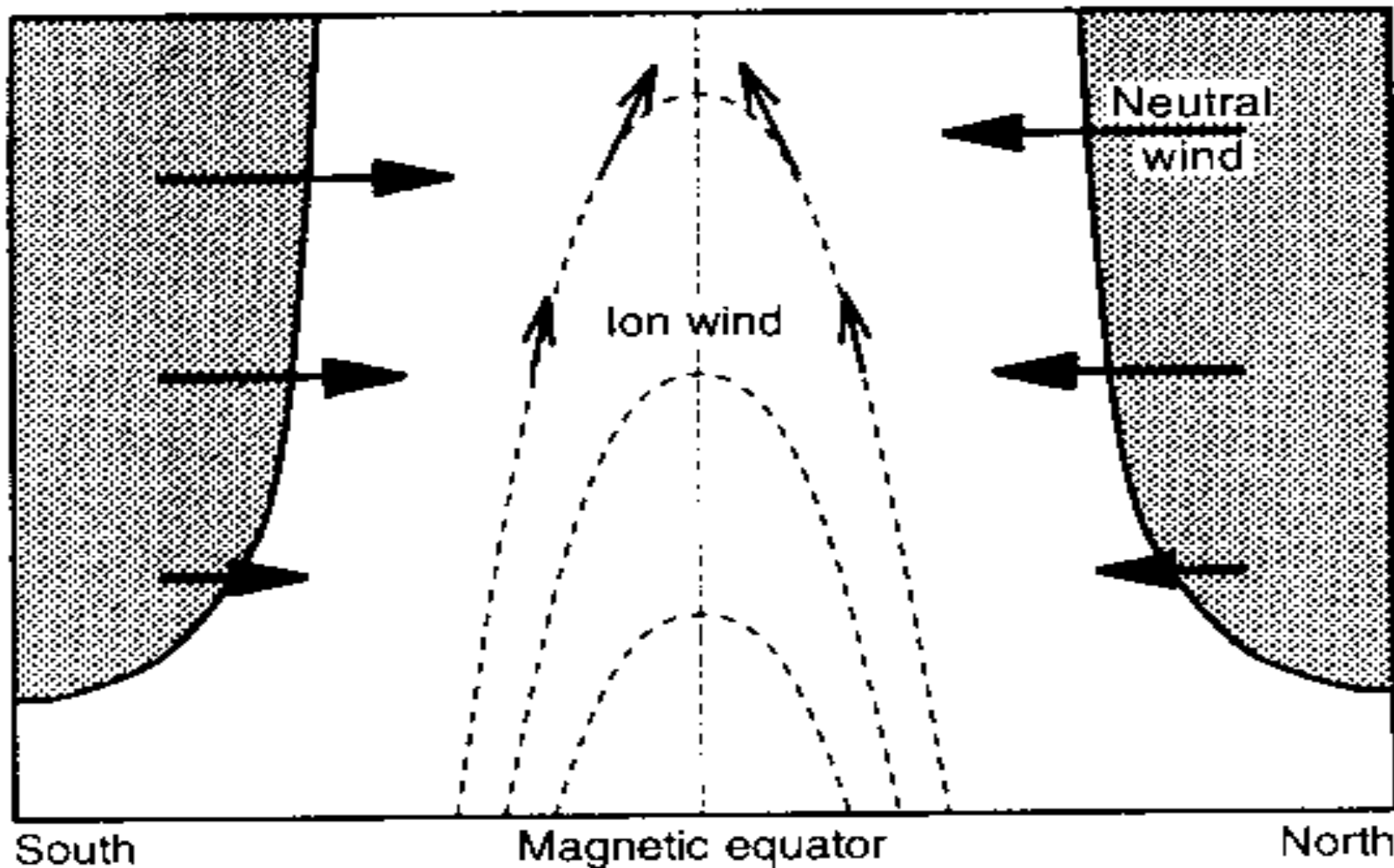
Observing TAD/TIDs

Motion of the thermosphere is notoriously difficult to image.

We rely on a proxy: the movement of the embedded ionosphere, which reflects and refracts radio waves. However...

...ions are constrained to spiral about field lines. Thus, horizontal motion of the thermosphere will result in a vertical component to the motion of the ionosphere (in the general case where $0 < \text{dip angle} < 90$)

TAD/TID coupling



hmF2 signature of TAD

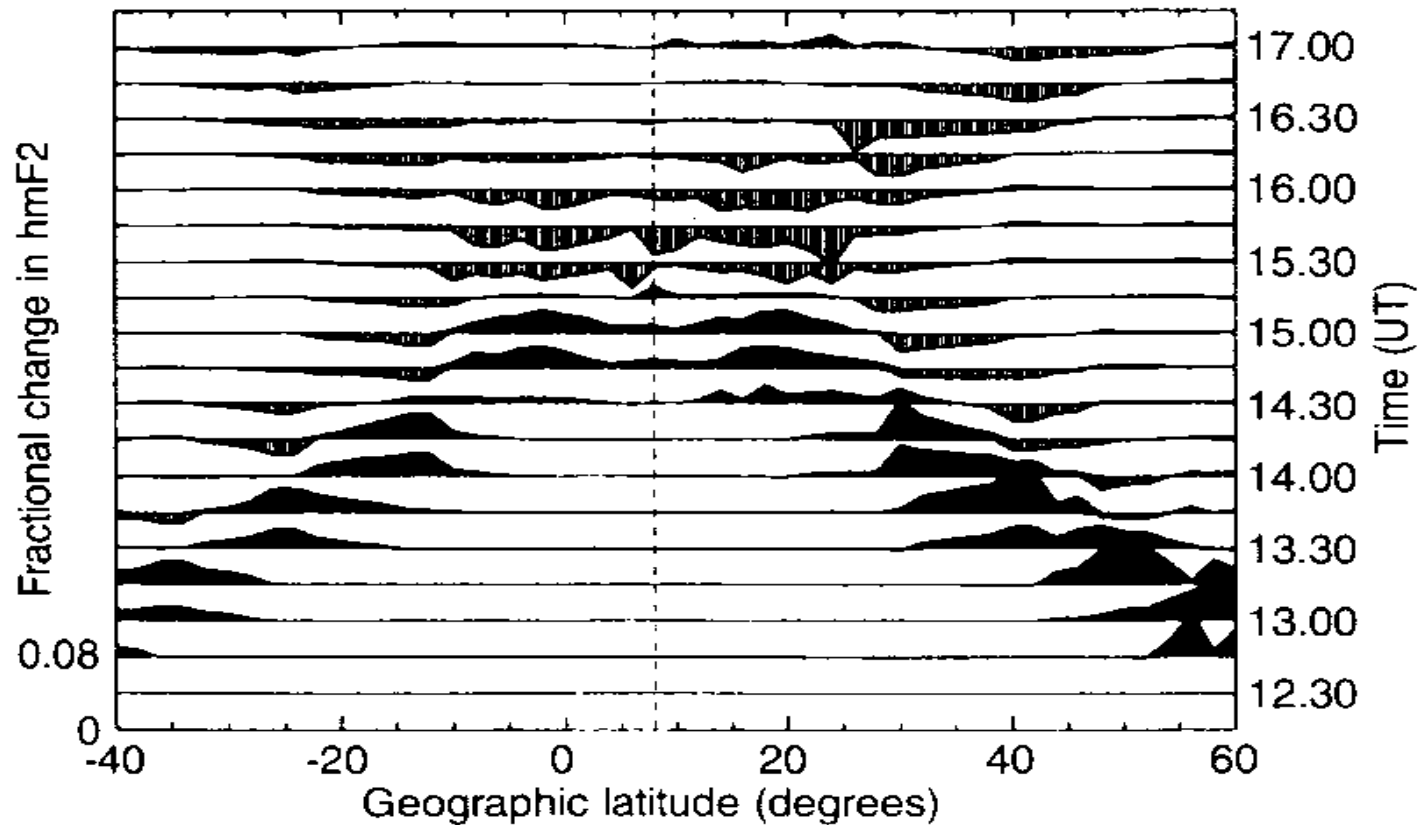


Fig. 4. Change in hmF2 (the height of the F2 peak), relative to steady-state conditions. The format of the plots is identical to that of Fig. 1

NmF2 signature of TAD

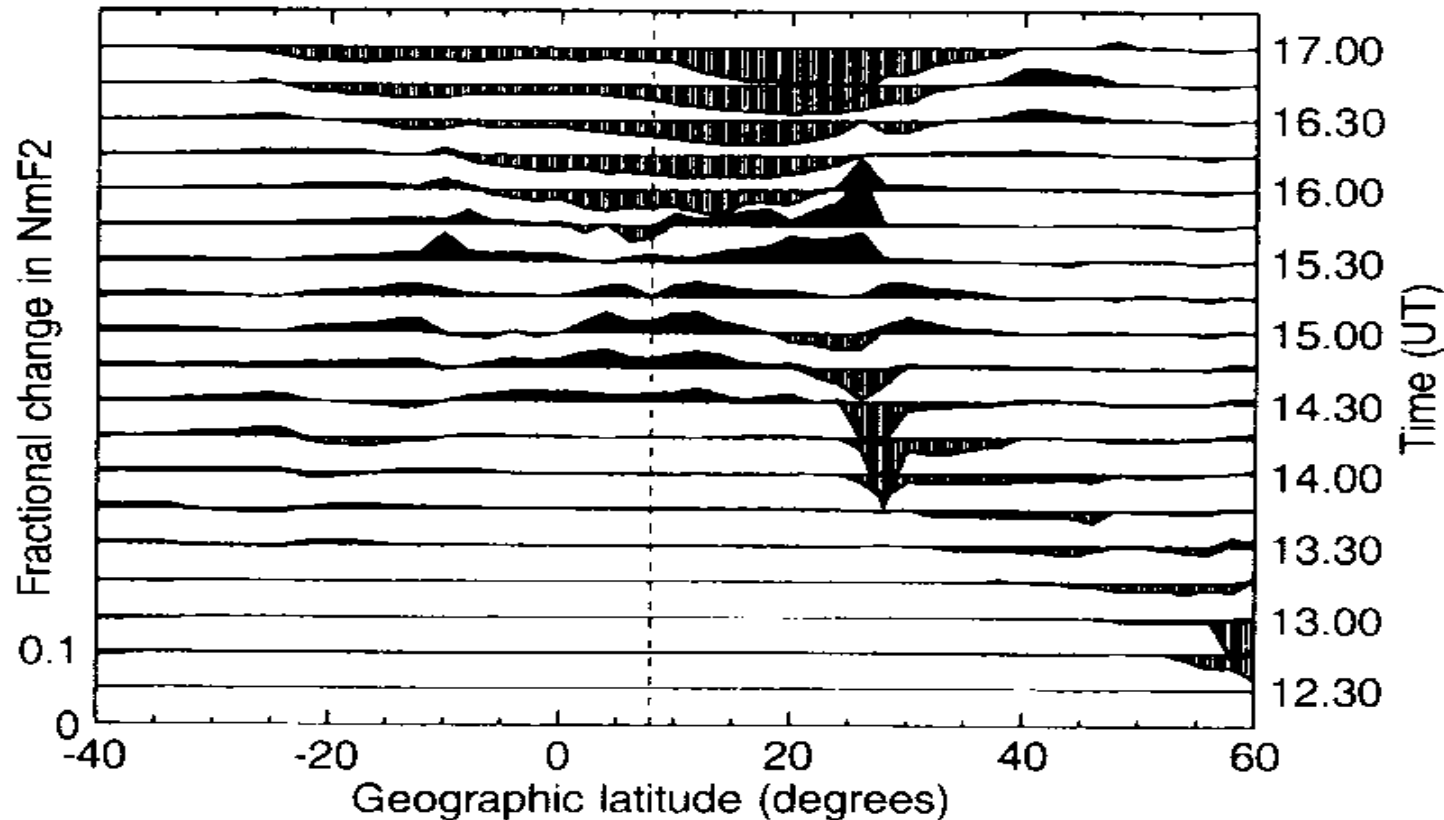
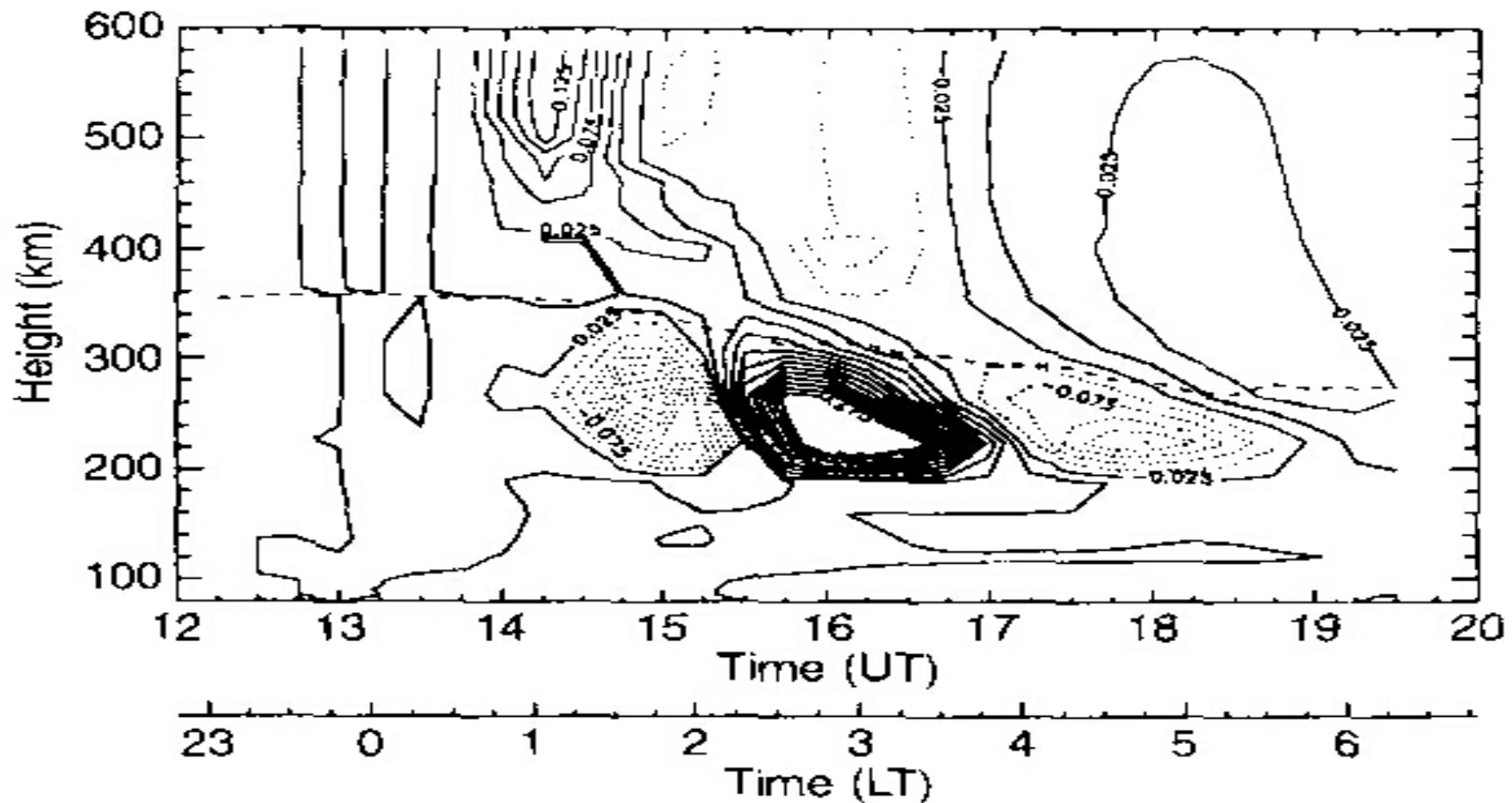


Fig. 5. Change in NmF2 (the density of the F2 peak), relative to steady-state conditions. The format of the plots is identical to that of Fig. 1

Ionospheric signature of TAD



Polar Vortices / Flywheel

Ions driven by a large electric field can accelerate to $\sim$ km/s

Lorentz forcing will impart momentum to the neutral gas

The polar thermosphere can thus 'spin up' into a flywheel, retaining its momentum, and driving the ionosphere long ($\sim$ hours) after the strong electric field conditions have vanished.

Super-rotation

Early satellite data showed that the equatorial thermosphere super-rotates in the evening sector

Average eastward zonal flow ~ 150 m/s
@ 350 km, 50 m/s @ 200 km, most pronounced @ 2100-2400 LT

Associated with solar heating

A light blue silhouette of a world map serves as a background for the title text.

Ionospheric features

Conductivities

- ◆ **Hall conductivity:** due to currents flowing perpendicular to the electric field
- ◆ **Pedersen conductivity:** due to currents flowing parallel to the electric field

F-region Motor

- ◆ During the daytime, the conductivity along the field lines is high
- ◆ The F-region plasma will thus be exposed to $\sim$ the same electric field which will tend to drive the F-region plasma like a motor with a velocity

$$\mathbf{v} = \frac{\mathbf{E} \times \mathbf{B}}{|\mathbf{B}|^2} \quad (6-20)$$

F-region dynamo

- ◆ In the F-region we have a conducting ‘slab’, defined at low altitudes by recombination erosion, and at high altitudes by the decreasing density gradient
- ◆ The vertical wind is always small
- ◆ The horizontal neutral wind, in equatorial regions, blows zonally (i.e., east-west)

F-region dynamo

Assume that the magnetic field lines are infinitely long and horizontal

In the reference frame of the neutral gas the **perturbed electric current** is

$$\mathbf{J}' = \sigma \cdot \mathbf{E}' \quad (6-21)$$

with conductivity, σ

In the Earth's reference frame, current is

$$\mathbf{J} = \sigma \cdot (\mathbf{E} + \mathbf{U} \times \mathbf{B}) \quad (6-22)$$

F-region dynamo

Therefore an electric current $\sim 10 \text{ nA/m}^2$ will flow upwards¹ due to the dynamo effect:

$$|\mathbf{J}| = \sigma_P u |\mathbf{B}| \quad (6-23)$$

¹ i.e., the electric field is down and positive ions move up while electrons move down

F-region dynamo

An electric field $\sim \text{mV m}^{-1}$ must build up in the vertical direction as a result of the charge separation to oppose the current flow.

$$|\mathbf{J}_z| = \sigma_P E_z + \sigma_P u B = 0 \quad (6-24)$$

Plasma drift speed = zonal neutral wind speed
 Night-time approximation of insulating slabs at ends
 (top and bottom of ionosphere)

E-region dynamo

- ◆ During the day, the ‘insulating ends’ approximation is not valid, and the current flow into the conducting E-region must be taken into account.
- ◆ Daytime E-region conductivity $>$ F-region conductivity (integrated up the field line), so the E-region ‘controls’ the dynamo process...

E-region dynamo (1)

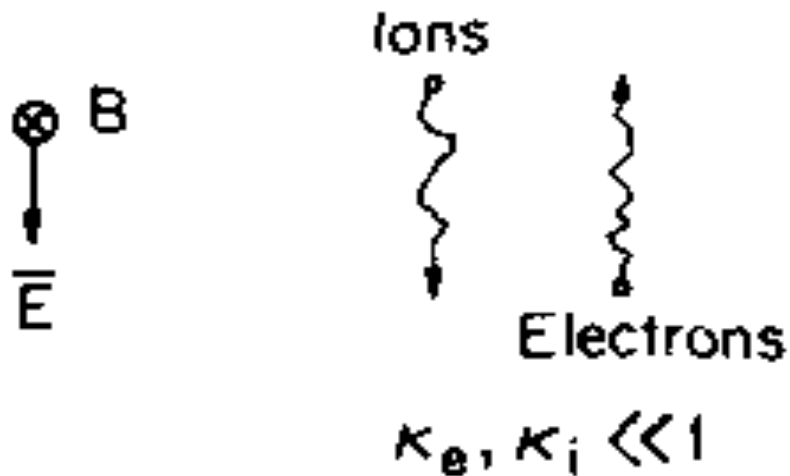
- ◆ In the E-region between ~ 85 km and ~ 140 km, the collision frequency is greater than the ion gyro frequency but less than the electron gyro frequency.
- ◆ Thus, ions move with the neutral wind while electrons move perpendicular to the neutral wind.
- ◆ An electric field is produced.

E-region dynamo (2)

- ◆ There will be an additional contribution to the electric current from the Hall conductivity (perpendicular to the winds)
- ◆ But no current can flow vertically, since the E-region is bounded above and below by ‘insulators’
- ◆ So the vertical Pedersen current must cancel the vertical Hall current.

Ion Gyrofrequency (ω) vs Collision Frequency (ν)

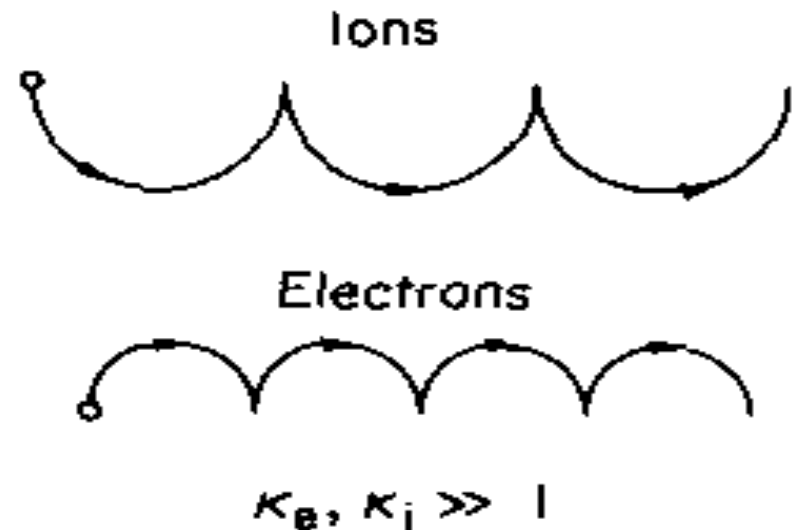
Collisional Case



$$\nu \gg \omega$$

Collision freq $\gg$ ion gyrofrequency

(b) Collisionless Case



$$\nu \ll \omega$$

Collision freq $\ll$ ion gyrofrequency

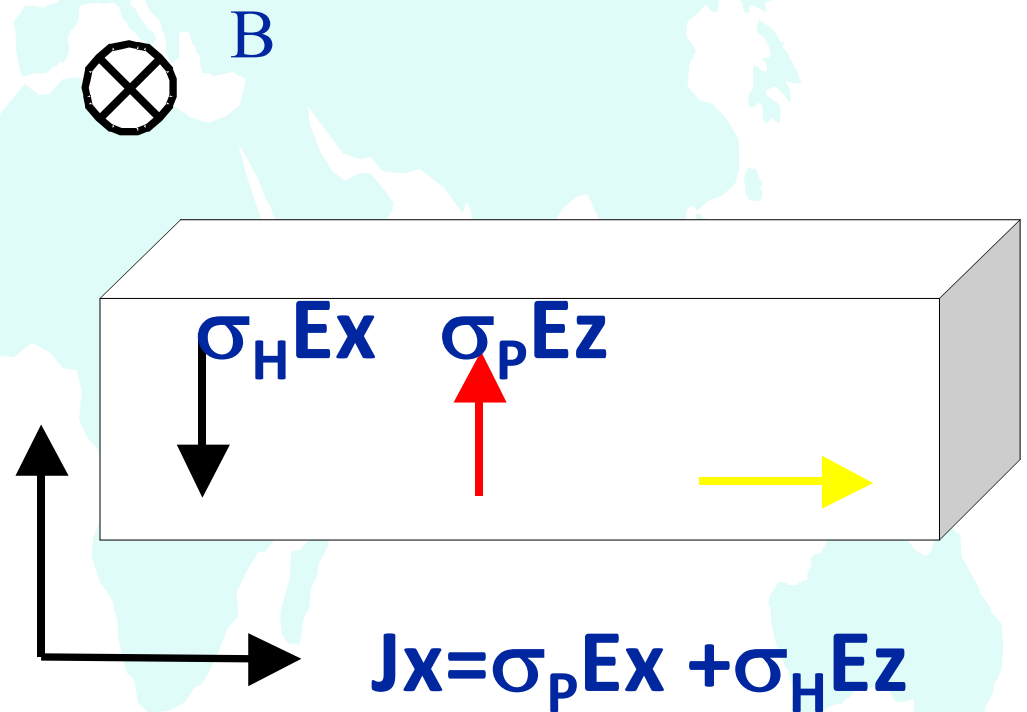
Conductivities help create the equatorial electrojet

$$\sigma_H E_x = \sigma_P E_z$$

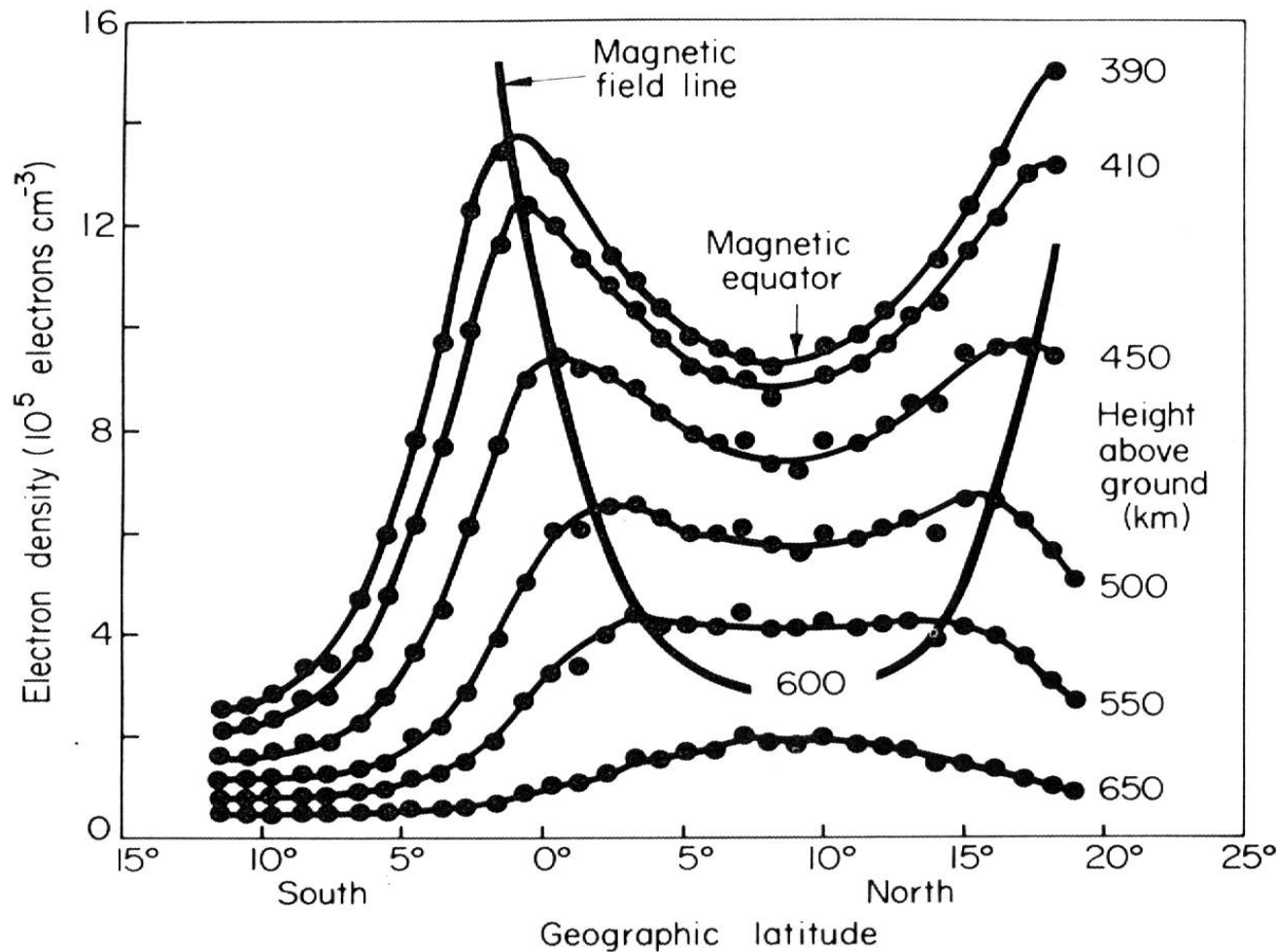
$$E_z = \left(\frac{\sigma_H}{\sigma_P} \right) E_x$$

$$J_x = \sigma_H E_z + \sigma_P E_x$$

$$J_x = \underbrace{\left(\frac{\sigma_H^2}{\sigma_P^2} + 1 \right)}_{\text{Cowling conductivity}} \sigma_P E_x$$



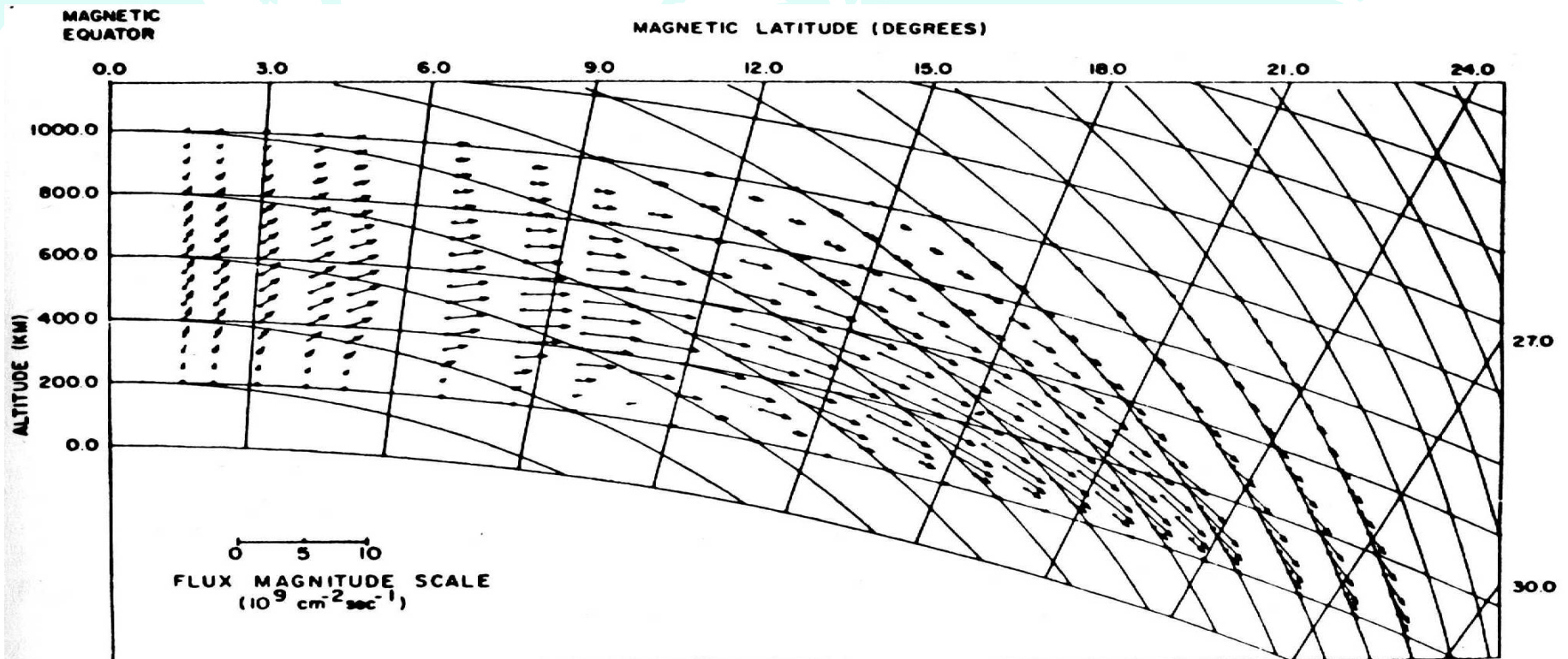
... and the Equatorial Anomaly



Equatorial Anomaly

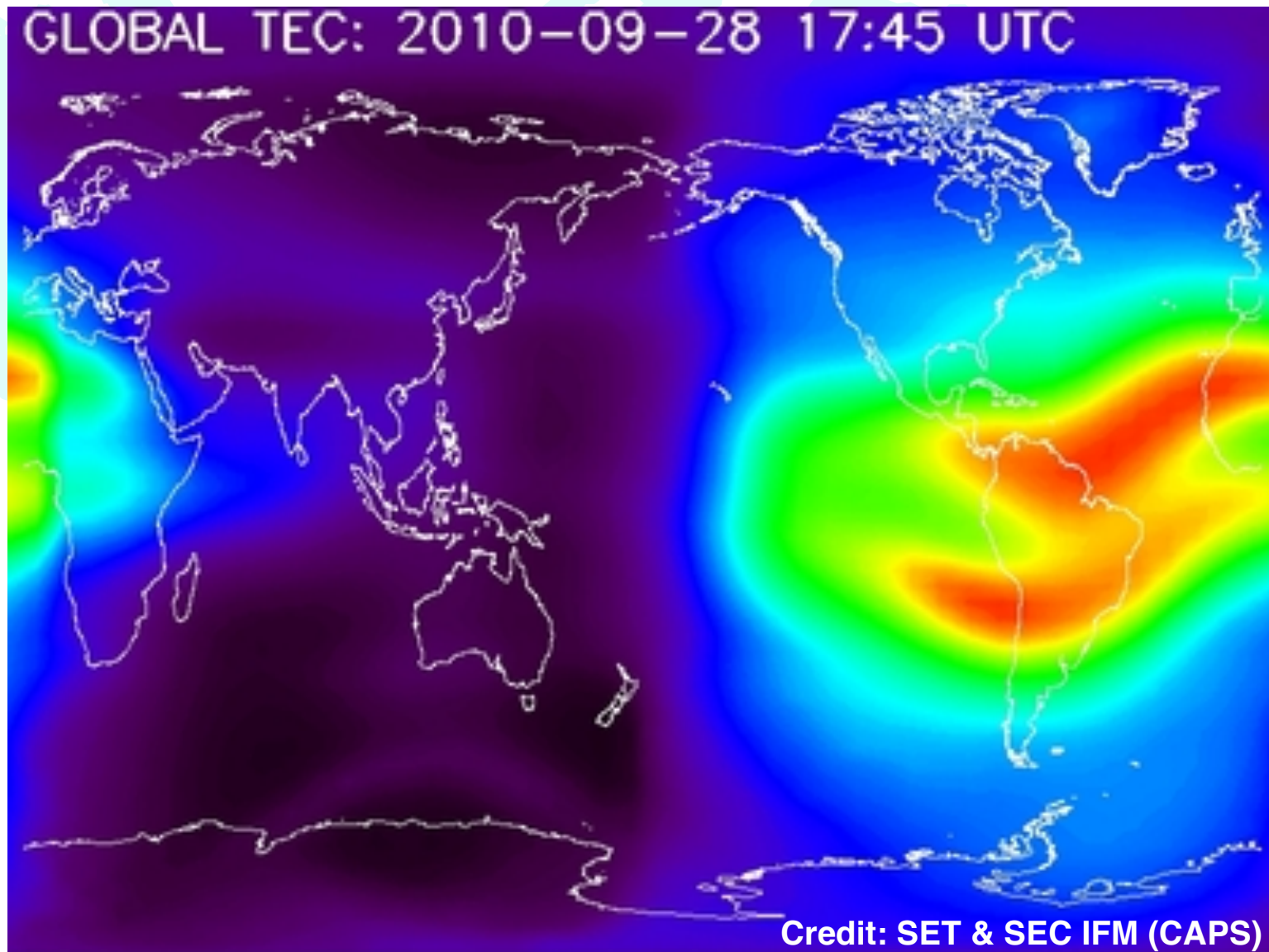
- ◆ aka Appleton Anomaly
- ◆ We might expect the electron density to maximize over the equator at equinox. However, it actually maximizes ~ 10 - 20 degrees N and S, with a small minima at the equator.
- ◆ To explain this, we invoke the ‘fountain effect’.

Fountain Effect



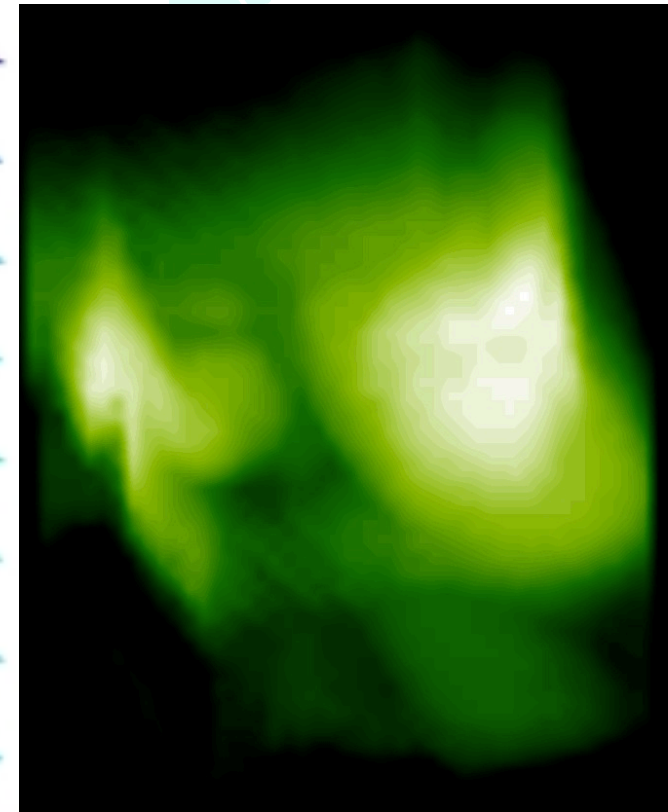
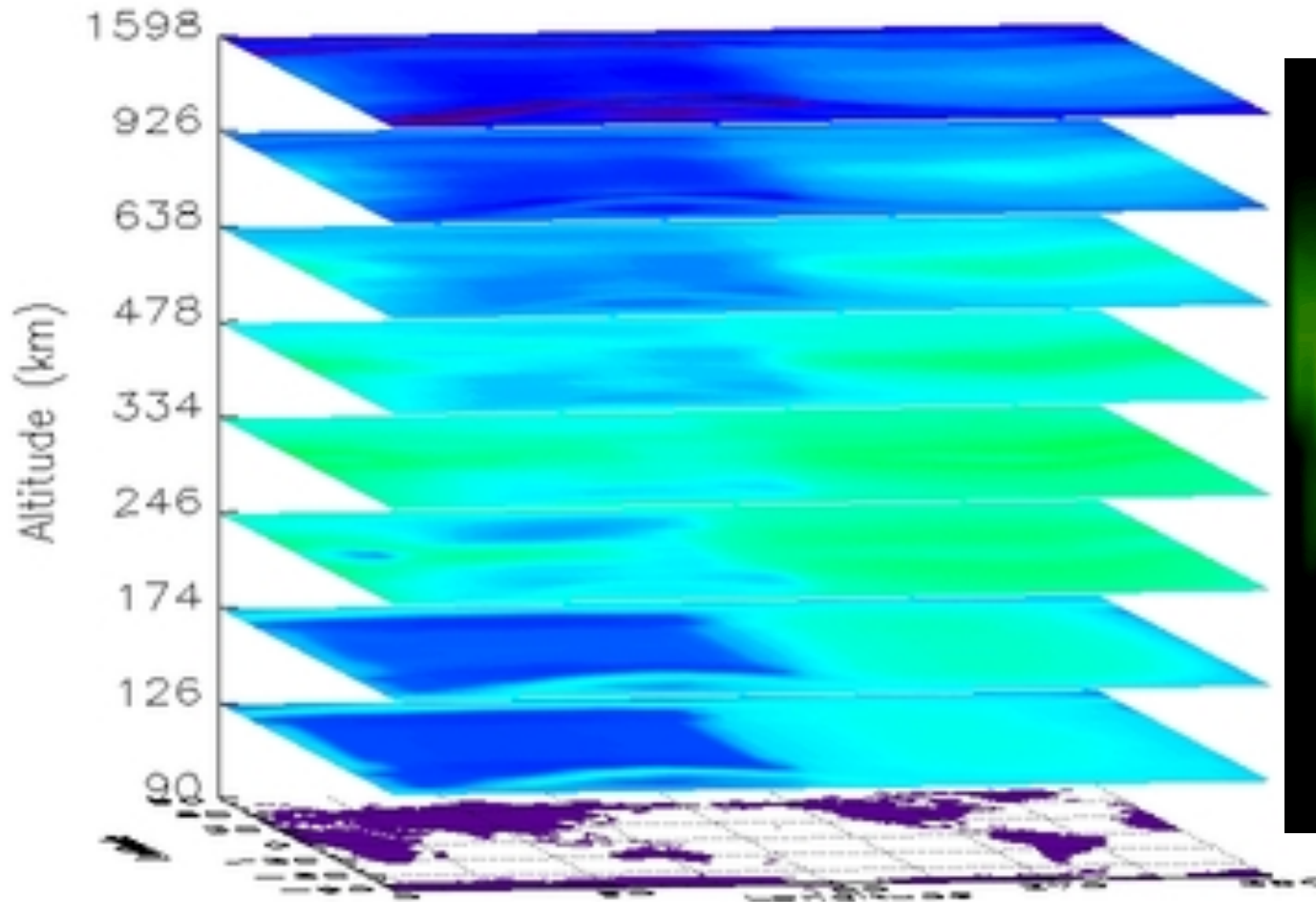
Equatorial Electrojet drives the F-region motor. E-field is zonal, magnetic field is meridional, so plasma drift is vertically upwards. Plasma then descends down the magnetic field lines either side.

Equatorial Anomaly



Equatorial Anomaly

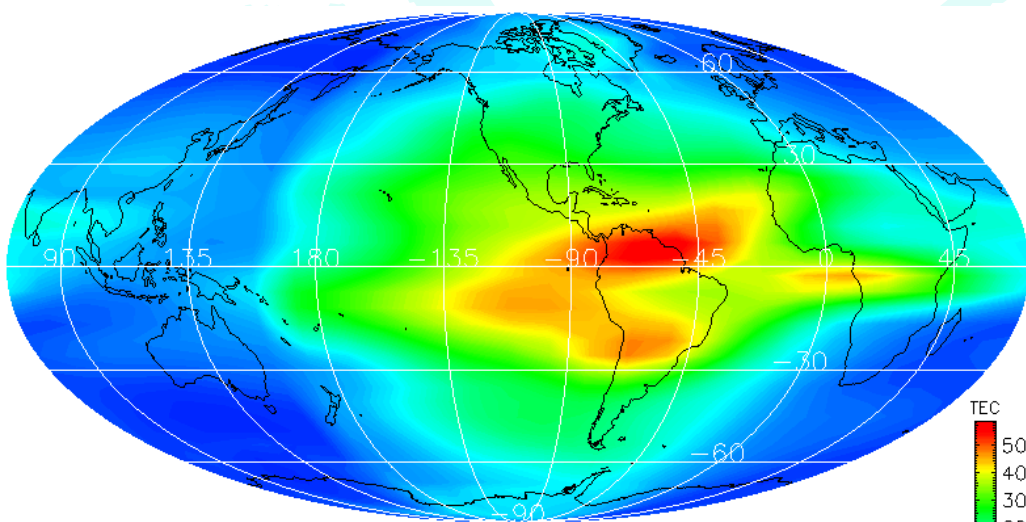
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electron densities

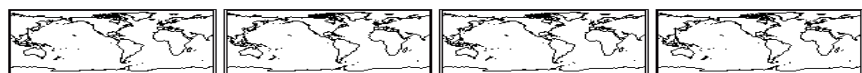
Credit: SET & SEC IFM (CAPS)

Equatorial Anomaly

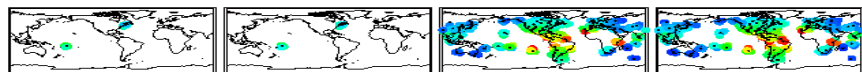


USU SWC 2011/09/27 19:15 UT

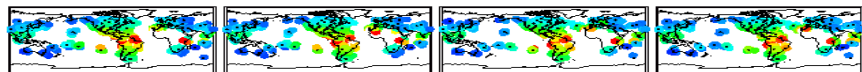
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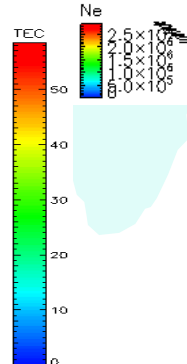
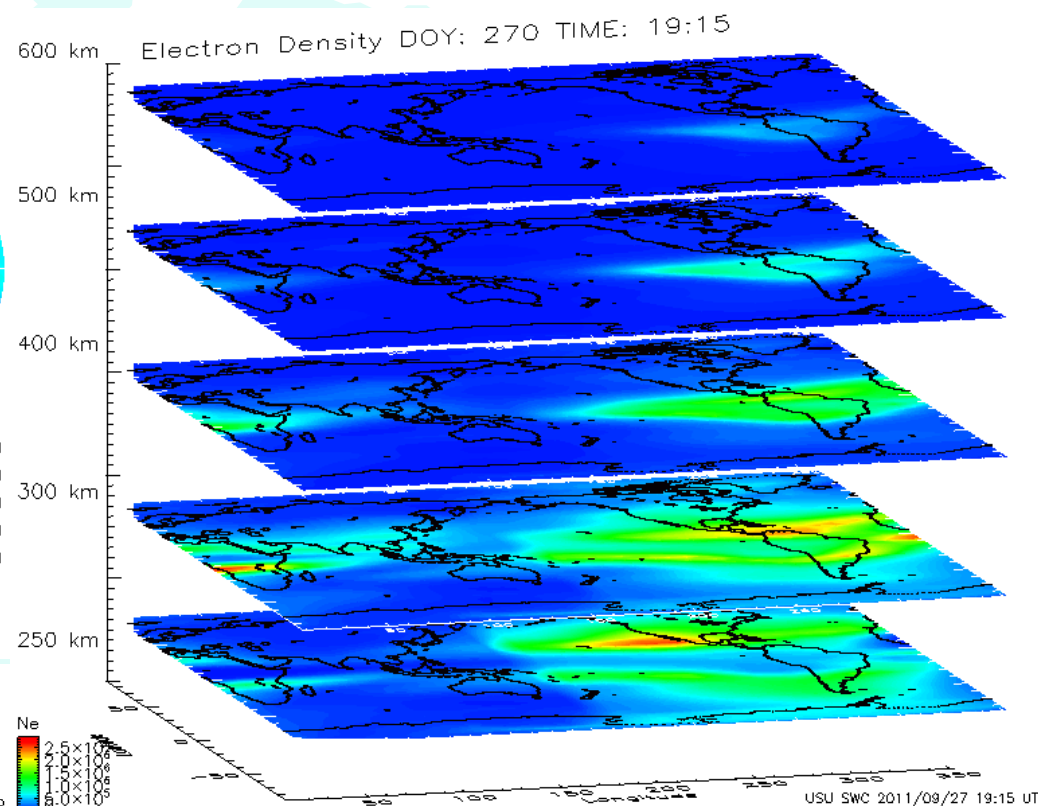


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<https://SpaceWx.com>

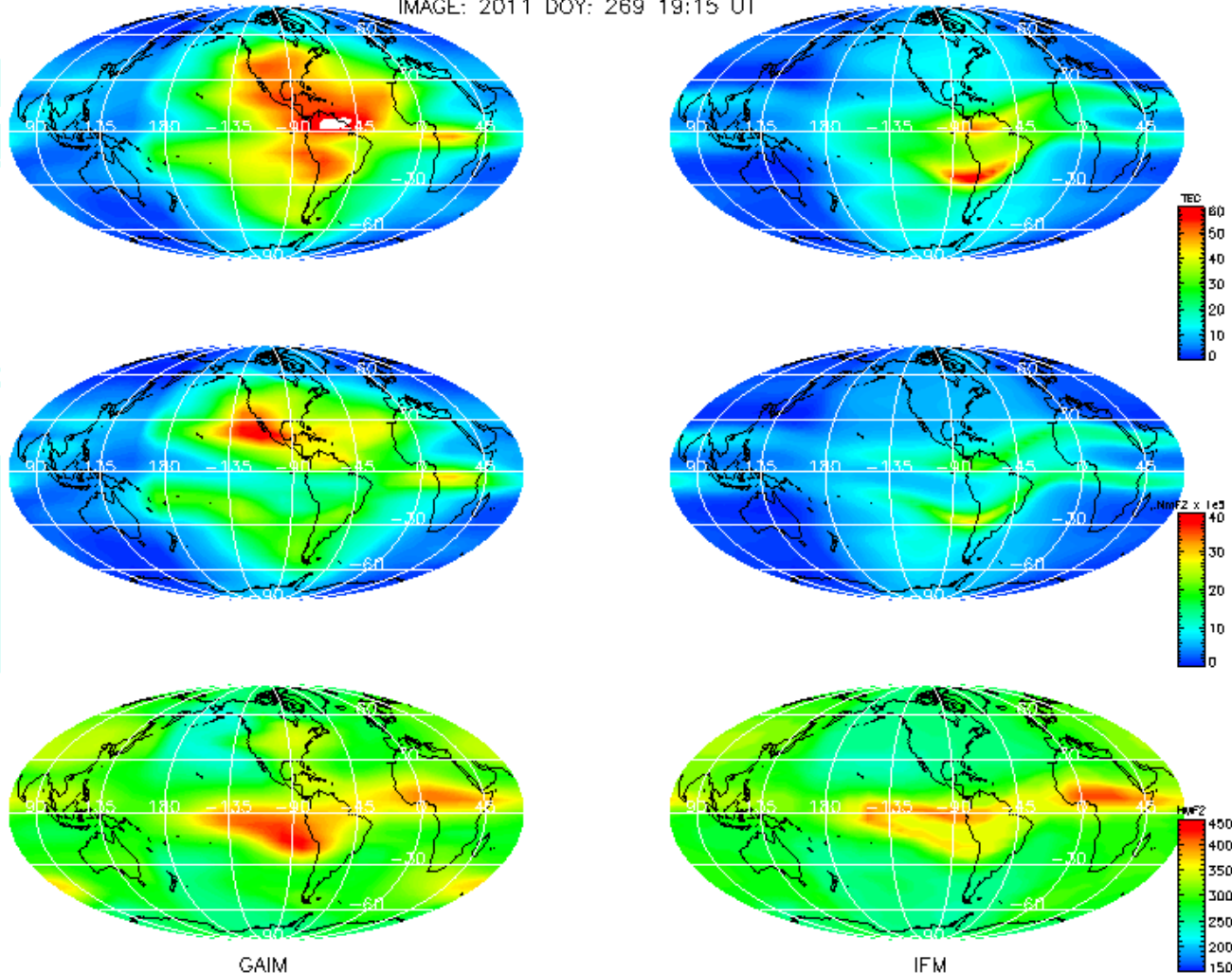


Credit: USU SWC

W. Kent Tobiska <ktobiska@spacewx.com>

Equatorial Anomaly

CURRENT: 2011 DOY: 270 19:00 UT
IMAGE: 2011 DOY: 269 19:15 UT



Credit: USU SWC

Resources

The Earth's Ionosphere, Kelley,
ISBN 0124040136, Academic Press
(*may be out of print*)

The Solar-Terrestrial Environment, Hargreaves,
ISBN 0521427371, Cambridge University Press

STEP Handbook of Ionospheric Models, ed. Robert
Schunk, Utah State University

Space Weather & Telecommunications, John M.
Goodman, ISBN 0-387-23670-8 (hc), Springer, 2005

Credit: Tobiska

Summary

✓ The planetary space environment - II

✓ Earth

✓ Ionospheric physics

✓ Thermospheric coupling

✓ Electron-ion production and loss processes (solar EUV photoionization, charged particle precipitation, Joule heating, waves, winds)

✓ Ionospheric Structure and Dynamics

✓ Ionospheric features